

REPORT NO. 2973P/A

STATOIL 30/3-3 NORWEGIAN NORTH SEA WELL:
BIOSTRATIGRAPHY OF THE INTERVAL 1900m - 3418m

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CONTENTS

	<u>Page No.</u>
I SUMMARY	1
II INTRODUCTION	3
III BIOSTRATIGRAPHIC SUCCESSION	5
IV LITHOSTRATIGRAPHIC SUCCESSION	7
V LITHOSTRATIGRAPHY	9
VI BIOSTRATIGRAPHY	15
(1) TERTIARY	15
(2) CRETACEOUS	18
(3) CRETACEOUS - JURASSIC	26
(4) JURASSIC	28
VII GEOLOGICAL HISTORY	50
VIII SELECTED REFERENCES	53

Appendices

Appendix 1	Lithological Descriptions of Sidewall Cores	55
Appendix 2	Lithological Descriptions of Core Pieces	59
Appendix 3	Rhaetian - Volgian Dinocyst Zonation	60

Figures

Figure 1	Legend	
Figures 2-10	Biostratigraphic Analysis Charts (Ditch Cuttings) Nos. 1-5	
Figures 11-13	Biostratigraphic Analysis Charts (Sidewall Cores and Cores) Nos. 1-2	

Enclosure

Enclosure 1	Summary Log (1 500 Scale)	
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SUMMARY

1. The youngest sediments analysed in this well are shales and tuffs of Late Palaeocene age, assignable to the Rogaland Group, Balder Formation.
2. Shales of the Sele Formation and the argillaceous/arenaceous Lista Formation underlie the Balder Formation and rest conformably upon the calcareous Montrose Group, Maureen Formation Equivalent of Early Palaeocene, Danian age.
3. The Early Palaeocene, Danian rests unconformably upon the argillaceous Shetland Group of Late Cretaceous, late Maastrichtian - Maastrichtian to Santonian age.
4. An unconformity separates the Shetland Group from the underlying Lower Cretaceous, Cromer Knoll Group of middle Albian - Barremian age. Sediments of this age are represented by a condensed sequence of claystones and limestones.
5. The Cromer Knoll Group rests unconformably upon an essentially argillaceous sequence of early Ryazanian - late Volgian to early Callovian age, assignable to the Humber Group, Kimmeridge Clay and Heather Formations.
6. Shales of the Heather Formation rest unconformably upon a thin development of the arenaceous Tarbert Formation of the Brent Group, of Middle Jurassic, early Bathonian - late Bajocian age.
7. A probable unconformity separates the Tarbert Formation from the underlying Ness Formation of middle - early Bajocian age.
8. The lithologically variable Ness Formation is conformably underlain by the predominantly arenaceous, Unassigned Unit of middle - early Bajocian to earliest Bajocian - middle Toarcian age.

9. The Unassigned Unit rests with apparent conformity upon the Dunlin Group, Drake Formation which rests upon an intra-Pliensbachian "Sand Unit" and a Pliensbachian - Sinemurian "Shale Unit".

10. The well terminated within the arenaceous Statfjord Formation of probable Sinemurian age.

INTRODUCTION

This report summarises the results of the micropalaeontological, palynological and stratigraphical analyses which have been carried out on material received from the section 1900m-3418m from the Statoil 30/3-3 Norwegian North Sea Well under Project No. RRPS/834/A/10581.

The following analyses were carried out:

Lithology: 300 ditch cuttings, 33 sidewall core and 4 core samples from the entire section.

Micropalaeontology: 146 ditch cuttings and 9 sidewall core samples from the interval 1900m-3380m.

Palynology: 64 ditch cuttings, 25 sidewall core and 4 core samples from the interval 2500m-3418m.

The basic breakdown obtained by these analyses has already been communicated by telex and telephone and forms the framework of factual information on which this report is based. A summary of the sequence penetrated in this well can be seen on pages 5 and 6.

All rock colour references in this report conform to the "Rock-Color Chart" distributed by the Geological Society of America. The lithological descriptions of the sidewall cores can be seen in Appendix 1, whilst those of the core samples occur in Appendix 2.

It should be noted that the lithostratigraphic terminology is taken from Deegan and Scull (1977). The stratigraphic significance of the Rhaetian - Volgian dinocyst zones is summarised in Appendix 3.

The prepared samples and recorded information are now filed and curated in the confidential records section of these laboratories.

We wish to acknowledge the continued co-operation and assistance received from the various members of Statoil with whom we have been associated during the course of this work.

Robertson Research staff involved in this study were:

Jim Fenton : Palynology and Project Co-ordinator

Paul Frame : Micropalaeontology

Nick Miles : Palynology

Alison Shaw : Lithostratigraphy

Dave Shipp : Micropalaeontology

III

BIOSTRATIGRAPHIC SUCCESSION

<u>Age</u>		<u>Tops</u>
Late Palaeocene		1900m(top not seen)
Early Palaeocene	Danian	2170m
-----Unconformity-----		
	(late Maastrichtian - Maastrichtian	2210m
	(
	(early Maastrichtian	2280m
	(
Late Cretaceous	(late Campanian	2340m
	(
	(early Campanian	2425m
	(
	(Santonian	2652.50m
-----Unconformity-----		
Early Cretaceous	middle Albian - Barremian	2655m
-----Unconformity-----		
Early Cretaceous - Late Jurassic	early Ryazanian - late Volgian	2664m(log)
	(middle - early Volgian	2740m
	(
	(Kimmeridgian	2770m(SWC)
	(
	(early Kimmeridgian - middle Oxfordian	2782.50m
Late Jurassic	(
	(early Oxfordian	2855m(SWC)
	(
	(late Callovian	2887.50m
	(
	(middle Callovian	2917.50m
	(
	(early Callovian	2970m
-----Unconformity-----		
	(early Bathonian - late Bajocian	2977.50m(log)
	(
Middle Jurassic	(-----Unconformity-----	
	(
	(middle - early Bajocian	2991m(log)

Middle - Early Jurassic	earliest Bajocian - middle Toarcian	3070m
Early Jurassic	(early Toarcian	3186m(SWC)
	(
	(late Pliensbachian - Sinemurian	3220m(SWC)- 3418m

This breakdown has been achieved by the analysis of ditch cuttings, sidewall core and core piece samples. Electric log information was made available.

LITHOSTRATIGRAPHIC SUCCESSION

<u>Units</u>	<u>Age</u>	<u>Tops</u>
Rogaland Group:		
Balder Formation	Late Palaeocene	1900m(top not seen)
Sele Formation	Late Palaeocene	2006m(log)
Lista Formation	Late Palaeocene	2070m
Montrose Group:		
Maureen Formation Equivalent	Early Palaeocene, Danian	2170m
Shetland Group:		
	late Maastrichtian - Maastrichtian to Santonian	2210m
Cromer Knoll Group:		
	middle Albian - Barremian	?2655m
Humber Group:		
Kimmeridge Clay Formation	early Ryazanian - late Volgian to Kimmeridgian	2664m(log)
Heather Formation	Kimmeridgian - early Calloviaian	2780m(log)
Brent Group:		
Tarbert Formation	early Bathonian - late Bajocian	2977.50m(log)
Ness Formation	middle - early Bajocian	2991m(log)
Unassigned Unit:		
	middle - early Bajocian to earliest Bajocian - middle Toarcian	3034m(log)

Dunlin Group:

Drake Formation	earliest Bajocian - middle Toarcian to late Pliensbachian - Sinemurian	3116m(log)
"Sand Unit"	late Pliensbachian - Sinemurian	3241m(log)
"Shale Unit"	late Pliensbachian - Sinemurian	3288m(log)
Statfjord Formation	late Pliensbachian - Sinemurian	3387m(log)- 3418m

LITHOSTRATIGRAPHY

The lithostratigraphic nomenclature used in this report is taken from Deegan and Scull (1977). Wireline logs were provided for use in the study and all but four of the lithostratigraphic boundaries are based on log breaks.

All depths quoted are uncorrected drillers depths unless otherwise stated.

LITHOSTRATIGRAPHIC UNITS

ROGALAND GROUP: 1900m(top not seen)-2170m; Age: Late Palaeocene.

Balder Formation: 1900m(top not seen)-2006m(log).

The predominant lithologies in this unit are shales which are light grey to greenish grey, greyish red and micromicaceous. The shales become mainly medium grey below 1980m. Subordinate tuffaceous shales are recorded throughout and become particularly common below 1985m; these sediments are grey and white, waxy and rubbly.

Pyrite and siderite are common accessories throughout and traces of hard, pale orange limestone are considered to represent stringers or concretions.

Sele Formation: 2006m(log)-2070m.

The upper boundary of this formation is based on a log break. The lithologies comprise shales which are weakly fissile, finely laminated and medium dark grey. Traces of limestone, siderite and pyrite occur throughout. Lithologies are masked by casing cement and lost circulation material between 2050m and 2065m.

Lista Formation: 2070m-2170m.

The upper limit of this unit is marked by an influx of fine to medium grained, loose sand in ditch cuttings samples. Shales remain the predominant lithology; these are medium dark grey to olive grey and greenish grey, micromicaceous and non-laminated. Traces of glauconite, siderite, pyrite and limestone are recorded. Fragments of limestone become particularly common below 2150m, consisting of hard, medium grey, slightly argillaceous limestone and probably

representing stringers or concretions.

MONTROSE GROUP: 2170m-2210m; Age: Early Palaeocene, Danian.

Maureen Formation Equivalent

The upper boundary of this formation is based on the increase in the calcareous nature of the shales. These shales are olive grey to medium grey and weakly calcareous, becoming medium light grey and moderately calcareous below 2185m. Traces of crystalline calcite and glauconite are recorded with subordinate, light brownish grey limestone below 2185m. Glauconite grains become abundant below 2205m.

SHETLAND GROUP: 2210m-?2655m; Age: Late Cretaceous, late Maastrichtian - Maastrichtian to Santonian.

The upper limit of this unit is based on an influx of white, slightly chalky limestone in the ditch cuttings samples at 2210m. Wireline log evidence suggests a depth of 2211m for the top of this limestone. This discrepancy is explained by the difference between drillers and loggers depths, which may be as much as 4m. Calcareous claystones are the predominant lithology below 2215m. These sediments are medium grey to greenish grey, olive grey and micromicaceous. Traces of glauconite, pyrite, calcite and common microfossils are recorded throughout. Subordinate limestone stringers occur sporadically, consisting of white to light grey limestone. A thicker development of pale brown limestone is indicated in wireline logs between 2412m(log) and 2424.50m(log).

Traces of fine, loose sand are recorded intermittently throughout, probably representing thin stringers of sandstone. Rare siderite occurs below 2590m and becomes abundant below 2650m.

CROMER KNOLL GROUP: ?2655m-2664m(log); Age: Early Cretaceous, middle Albian - Barremian.

The upper boundary of this unit is tentatively placed at 2655m, based on biostratigraphic evidence. No change in lithotype is observed until a depth of 2560m, where fragments of white to very light grey limestone are recorded, probably representing the older part of this interval.

Calcareous claystones are the predominant lithology; these sediments are medium grey to medium dark grey with traces of calcite, pyrite and sand. Minor white to very light grey limestones occur at 2659m(log) and between 2662m(log) and 2664m(log).

HUMBER GROUP: 2664m(log)-2977.50m(log); Age: Early Cretaceous - Late Jurassic, early Ryazanian - late Volgian to Late Jurassic, early Callovian.

Kimmeridge Clay Formation: 2664m(log)-2780m(log); Age: Early Cretaceous - Late Jurassic, early Ryazanian - late Volgian to Late Jurassic, Kimmeridgian.

The upper boundary of this formation is based on a log break reflecting a change in lithotype to the black shales typical of the Kimmeridge Clay Formation. Due to the discrepancy between drillers and loggers depths, however, these shales are first recorded in the ditch cuttings sample at 2662.50m.

The shales are fissile, olive black to brownish black, dark grey, micromicaceous and non-calcareous. Traces of pyrite, siderite, carbonaceous material and sand occur sporadically throughout. Increasing amounts of medium grey shale with subordinate brownish grey limestone are recorded below 2755m, with abundant sphaerosiderite between 2755m and 2765m.

Heather Formation: 2780m(log)-2977.50m(log); Age: Late Jurassic, Kimmeridgian to early Callovian.

The upper limit of this unit is based on a log break marking the top of a sequence of shales and minor limestones with a lower gamma ray response than that of the overlying Kimmeridge Clay Formation shales.

The shales are fissile, medium grey to dark grey, olive black, micromicaceous and locally highly micaceous. Subordinate limestones are brownish grey with calcite, probably representing stringers or concretions. Fragments of limestone become increasingly common in ditch cuttings samples below 2840m. Shale is present in subordinate quantities down to 2830m, below which depth shales become the predominant lithology. These shales are brownish grey to dark grey, olive black and micaceous, locally becoming slightly silty and dusky yellowish brown.

Pyrite and calcite are common throughout and abundant sphaerosiderite is recorded at 2890m and from 2910m to 2935m.

BRENT GROUP: 2977.50m(log)-3034m(log); Age: Middle Jurassic, early Bathonian - late Bajocian to middle - early Bajocian.

Tarbert Formation: 2977.50m(log)-2991m(log); Age: Middle Jurassic, early Bathonian - late Bajocian.

This formation comprises argillaceous sandstone, which is well cemented, medium light grey, very fine to fine grained and micaceous with carbonaceous fragments. Argillaceous partings of medium dark grey, micaceous shale are present locally.

Ness Formation: 2991m(log)-3034m(log); Age: Middle Jurassic, middle - early Bajocian.

This unit comprises a sequence of sandstones interbedded with shales, claystones, coals and argillaceous sandstones.

Between 2991m(log) and 3000m(log), brownish grey and olive black claystones predominate, with minor coals and sandstones. From 3000m(log) to 3017m(log), sandstones become the main lithology and these sediments are white, fine to medium grained and kaolinitic with rare coarse grains. Shales and coals dominate the section from 3017m(log) to 3028m(log) with sandstones and shales comprising the remainder of the interval. A coal marks the base of this unit.

UNASSIGNED UNIT: 3034m(log)-3116m(log); Age: Middle Jurassic, middle - early Bajocian to Middle - Early Jurassic, earliest Bajocian - middle Toarcian.

This interval has a similar log character to the sand unit which is usually assigned to the Etive/Rannoch Formations in this area of the North Sea. Biostratigraphic evidence suggests, however, that the lower part of this sand is older than nearby Brent Group sediments, hence this unit has not been assigned to the Brent Group.

The predominant lithology is sandstone which is white to pinkish grey, fine to coarse grained, kaolinitic, micaceous and locally weakly calcareous. Minor argillaceous interbeds are indicated on wireline logs, consisting of shales which are brownish grey to dark grey and micromicaceous with rare carbonaceous fragments.

DUNLIN GROUP: 3116m(log)-3387m(log); Age: Middle - Early Jurassic, earliest Bajocian - middle Toarcian to Early Jurassic, late Pliensbachian - Sinemurian.

One formation and two informal units are recognised within this group.

Drake Formation: 3116m(log)-3241m(log); Age: Middle - Early Jurassic earliest Bajocian - middle Toarcian to Early Jurassic, late Pliensbachian - Sinemurian.

This formation consists predominantly of shales which are dark yellowish brown to dark grey and olive black to olive grey, micromicaceous, locally silty with rare carbonaceous fragments and pyrite. A minor sandstone unit is recorded between 3152m(log) and 3166m(log). These sediments are friable, light grey to white, very fine to medium grained and kaolinitic with traces of pyrite.

"Sand Unit": 3241m(log)-3288m(log); Age: Early Jurassic, late Pliensbachian - Sinemurian.

This unit comprises sandstones which become finer grained and increasingly argillaceous with depth. The sandstones are white, fine grained and kaolinitic, locally with calcareous cement, becoming very pale orange to dark yellowish brown and very fine grained with minor shale partings.

"Shale Unit": 3288m(log)-3387m(log); Age: Early Jurassic, late Pliensbachian - Sinemurian.

The predominant lithology in this unit is shale which is olive black, micaceous and locally very weakly calcareous with carbonaceous fragments. Below 3346m(SWC), the shales become dusky yellowish brown micaceous and locally sandy, grading into argillaceous sandstone. Rare fragments of limestone and calcite are recorded between 3330m and 3345m, probably representing stringers or concretions.

Traces of white, fine grained sandstone occur below 3376m(SWC).

STATFJORD FORMATION: 3387m(log)-3418m; Age: Early Jurassic, late Pliensbachian - Sinemurian.

The upper boundary of this formation is marked by an influx of fine to coarse grained sand. This is first recorded in the ditch cuttings sample at 3380m, however, wireline logs indicate an upper limit of 3387m. This discrepancy is explained by the difference between drillers and loggers depths which may be up to 7m.

The sandstones are fairly friable, pinkish grey to white, fine to coarse grained and kaolinitic with carbonaceous fragments and traces of pyrite. These sediments are locally light grey and slightly argillaceous.

BIOSTRATIGRAPHYVI (1) TERTIARY

INTERVAL 1900m-2170m; LATE PALAEOCENE (top not seen)

Lithostratigraphic Units:

Rogaland Group;

Balder Formation: 1900m-2006m(log),

Sele Formation: 2006m(log)-2070m,

Lista Formation: 2070m-2170m.

Environment: marine, outer shelf to upper bathyal.

The age of this interval is based on the following lithostratigraphic and micropalaeontological criteria:

- the presence of volcanic, tuffaceous deposits from 1900m.
- the occurrence of the diatom *Coscinodiscus* sp.1 at 1920m and the foraminifer *Trochammina* sp.A at 2100m.

MICROPALAEONTOLOGY

The microfaunas of this interval are dominated by agglutinated species, particularly of the *Bathysiphon/Rhizammina* Group, *Haplophragmoides* spp., *Recurvoides* spp. and *Cyclammina* spp.

Common, red stained planktonic foraminifera including *Globigerina linaperta* and *G. triloculinoidea* were recovered from the uppermost section of this interval (1900m-1920m). The presence of the latter species and the associated red staining characteristically indicate the presence of the *Globigerina triloculinoidea* Assemblage of Early Eocene age. In view of the occurrence of typical Late Palaeocene tuffaceous deposits at 1900m, however, a Late Palaeocene age can be assigned to the interval and the above species are considered to be caved.

The first micropalaeontological evidence confirming the age of the interval occurs at 1920m with the presence of the diatom *Coscinodiscus* sp.1 indicating the development of the *Coscinodiscus* sp.1 Assemblage of Late Palaeocene age. The subsequent occurrence of *Trochammina* sp.A at 2100m further supports this age determination.

Early Palaeocene, Danian reworking is recorded at 1940m and 2020m.

INTERVAL 2170m-2210m; EARLY PALAEOCENE, DANIAN

Lithostratigraphic Unit:

Montrose Group;

Maureen Formation Equivalent.

Environment: marine, inner to outer shelf.

The upper limit and age of this interval are based on the following micropalaeontological criteria:

- the presence of *Globigerina trivialis* at and below 2170m.
- the subsequent occurrence of *Globorotalia pseudobulloides* (2180m) and *G. cf. compressa* (2190m).

The base of this interval is marked by an unconformity separating the Tertiary from the Late Cretaceous and resulting in the absence of part of the Early Palaeocene, Danian.

MICROPALAEONTOLOGY

As in the overlying section agglutinated foraminifera dominate the microfaunas of this interval. The presence of rare planktonic species, however, is also recorded.

The occurrence of *Globigerina trivialis* at 2170m is taken to represent the Late/Early Palaeocene boundary. The incoming of this species and the subsequent occurrence of *Globorotalia pseudobulloides* (2180m) and *G. cf. compressa* (2190m) also indicate penetration of the P.2 Zone of the Early Palaeocene, Danian. No evidence for the older P.1 Zone is recorded.

VI (2) CRETACEOUS

INTERVAL 2210m-2280m; LATE CRETACEOUS, LATE MAASTRICHTIAN - MAASTRICHTIAN

Lithostratigraphic Unit:
Shetland Group (part).

Environment: marine, inner to outer shelf.

The upper limit and age of this interval are based on the following micropalaeontological criterion:

- the appearance of *Pseudotextularia elegans fructicosa*, *P. elegans elegans* and *Globotruncana contusa* at 2210m.

MICROPALAEONTOLOGY

Planktonic foraminifera dominate the microfaunas of this interval, being represented mainly by the *Rugoglobigerina*/*Archaeoglobigerina* complex, *Globigerinelloides asper* and at the top of the interval, *Pseudotextularia* spp. and *Globotruncana contusa*.

The presence of common *Pseudotextularia elegans fructicosa*, *P. elegans elegans* and *Globotruncana contusa* at 2210m indicates a late Maastrichtian age at that depth. It has not proved possible to subdivide the late Maastrichtian and Maastrichtian deposits, although the assemblages recorded suggest that both intervals are present.

A number of other Late Cretaceous species also have their first occurrences in this interval. These include *Stensioina pommerana* (2210m), *Abathomphalus mayaroensis* (2225m (SWC)), *Reussella szajnochae szajnochae* (2230m), *Heterohelix costulata* (2230m), *Heterohelix globulosa* (2250m) and *Reussella szajnochae elongata* (2260m).

Tertiary caving occurs in this interval and continues through most of the underlying Late Cretaceous.

Late Cretaceous, Campanian reworking was recorded at 2240m, 2250m and 2260m.

INTERVAL 2280m-2340m; LATE CRETACEOUS, EARLY MAASTRICHTIAN

Lithostratigraphic Unit:
Shetland Group (part).

Environment: marine, inner to outer shelf.

The upper limit and age of this interval are based on the following micropalaeontological criterion:

- the influx of abundant *Rugoglobigerina/Archaeoglobigerina* complex together with common, consistently occurring *Rugoglobigerina rugosa rotundata* at 2280m.

MICROPALAEONTOLOGY

Planktonic foraminifera again dominate the microfaunas, being represented principally by the *Rugoglobigerina/Archaeoglobigerina* complex.

The marked influx of abundant *Rugoglobigerina/Archaeoglobigerina* complex at 2280m indicates the development of the *Rugoglobigerina/Archaeoglobigerina* Assemblage of early Maastrichtian age. The associated occurrence of common and consistently occurring *Rugoglobigerina rugosa rotundata* at and below this depth supports the age determination.

INTERVAL 2340m-2425m; LATE CRETACEOUS, LATE CAMPANIAN

Lithostratigraphic Unit:

Shetland Group (part).

Environment: marine, inner to outer shelf.

The upper limit and age of this interval are based on the following micropalaeontological criterion:

- the occurrence, at 2340m, of *Globotruncana* sp.1 and *G. marginata*.

MICROPALAEONTOLOGY

Planktonic foraminifera, principally of the *Rugoglobigerina/Archaeoglobigerina* complex, are again dominant throughout this interval.

The presence of *Globotruncana* sp.1 and *Globotruncana marginata* from 2340m indicates a late Campanian age for the interval. The occurrence of these two species allows the microfaunas of the interval to be assigned to the *Globotruncana marginata* Assemblage.

Tertiary caving is present throughout the section.

INTERVAL 2425m-2652.50m; LATE CRETACEOUS, EARLY CAMPANIAN

Lithostratigraphic Unit:

Shetland Group (part).

Environment: marine, inner to outer shelf.

The upper limit and age of this interval are based on the following micropalaeontological and palynological criteria:

- the occurrence of *Tritaxia dubia* and *Dorothia* sp.1 at 2425m.
- the subsequent occurrence of *Tritaxia tricarinata* (2535m).
- the appearance of *Callaiosphaeridium asymmetricum* at 2560m.

MICROPALAEONTOLOGY

Very consistent microfaunas, again dominated by the planktonic *Rugoglobigerina/Archaeoglobigerina* complex, occur throughout this interval. The appearance of the agglutinating species *Tritaxia dubia* and *Dorothia* sp.1 at 2425m, however, indicates that the early Campanian has been penetrated. The subsequent occurrence of *Tritaxia tricarinata* at 2535m supports this determination.

PALYNOLOGY

Assemblages recovered at and below 2500m are dominated by microplankton, with miospores being a very minor constituent. Palynofloras are characterised by the common to abundant occurrences of the *Spiniferites ramosus* group, *Ceratiopsis diebelii*, *Palaeoperidinium pyrophorum*, *Hystrichosphaeridium tubiferum*, *Areoligera* spp., *Chatangiella* spp./*Isabelidinium* spp., *Cyclonephelium* spp., *Chatangiella granulifera* and *Trithyrodinium suspectum*. Palaeocene cavings are conspicuous in most assemblages.

The overall composition of the palynofloras is indicative of a Campanian age, with the co-occurrences of large numbers of *C. diebelii*, *Spongodinium delitiense* and *Raphidodinium fucatum*, in association with taxa such as *T. suspectum*, *Alterbia acutula* (2500m), *Odontochitina* spp., *Chatangiella granulifera* and *Trichodinium castaneum*. An age no younger than early Campanian is indicated at 2560m by the appearance of *Callaiosphaeridium asymmetricum*.

The sidewall core at 2640m yielded a palynoflora of early Campanian age, characterised by the simultaneous presence of Callaiosphaeridium asymmetricum, Spinidinium clavum, Trigonopyxidia ginella, Hystrichosphaeridium sp.2 McIntyre, abundant specimens of T. suspectum and numerous specimens of Raphidodinium fucatum.

Reworked miospores are occasionally encountered reflecting recycling of deposits of Permo-Triassic and Late - Middle Jurassic ages.

INTERVAL 2652.50m-2655m; LATE CRETACEOUS, SANTONIAN

Lithostratigraphic Unit:
Shetland Group (part).

Environment: marine, inner to outer shelf.

The upper limit and age of this interval are based on the following micropalaeontological criterion:

- the occurrence of common, green stained spongodiscid radiolaria at 2625.50m.

The base of this interval is marked by an unconformity separating the Late and Early Cretaceous, resulting in the absence of any Coniacian to late Albian deposits.

MICROPALAEONTOLOGY

Moderately common foraminifera were recorded in this interval with planktonic and agglutinated species predominating.

No foraminiferal marker species were recovered from the interval. The presence of common, green stained spongodiscid radiolaria at 2652.50m, however, suggests the presence of Santonian deposits, these radiolaria being a common feature of the Santonian.

INTERVAL 2655m-2664m(log): EARLY CRETACEOUS, MIDDLE ALBIAN - BARREMIAN

Lithostratigraphic Unit:

Cromer Knoll Group.

Environment: shallow marine, inner shelf.

The upper limit and age of this interval are defined upon the following micropalaeontological and palynological criteria:

- the presence of *Praebulimina* sp.1 at 2655m and the overall nature of the microfaunas.
- the occurrence of a specimen of *Cribroridinium cornutum* at 2655m.

An unconformity is present at the base of this interval resulting in the absence of sediments of Hauterivian - late Ryazanian age.

MICROPALAEONTOLOGY

A marked decline in planktonic abundance occurs throughout this interval with an associated increase in agglutinated and, to a lesser extent, calcareous benthonic abundances.

The foraminifera recovered from this interval suggest that sediments of a broad age range are represented. The occurrence of *Praebulimina* sp.1 at 2655m indicates that sediments of middle Albian - Aptian age have been penetrated. This is supported by the predominance of agglutinated species in the samples, including *Recurvoides* spp. and *Glomospira* spp. The increased occurrence lower in the section of calcareous benthonic species, particularly *Lenticulina* spp. and ostracods, however, may also indicate the presence of sediments of an older, Barremian age. It is not possible, due to the sample intervals received, to subdivide this interval further or to determine whether it represents a condensed sequence of sediments or if there are any hiatuses present in the section. Consequently, only a broad middle Albian to Barremian age can be assigned.

PALYNOLOGY

Palynofloras recovered from 2655m and 2660m are considerably modified by caved taxa of Eocene/Palaeocene and Late Cretaceous ages. Assemblages are

characterised by common/abundant specimens of *Palaeoperidinium pyrophorum* and species of *Chatangiella*, associated with taxa such as *Raphidodinium fucatum*, *Odontochitina* spp., *Trithyrodinium suspectum* and *Spongodinium delitiense*. Miospores are a minor constituent. In situ palynomorphs are impoverished.

A single specimen of *Cribooperidinium cornutum* was recorded at 2655m, a taxon stratigraphically restricted to deposits of Barremian age. The only specimens thought to be in situ at 2660m are those of *Oligosphaeridium* complex exhibiting "typical" Early Cretaceous process development.

VI (3) CRETACEOUS - JURASSIC

INTERVAL 2664m(log)-2740m; EARLY CRETACEOUS - LATE JURASSIC, EARLY RYAZANIAN - LATE VOLGIAN

Lithostratigraphic Unit:

Humber Group (part);

 Kimmeridge Clay Formation (part).

Environment: marine, inner shelf with anaerobic/dysaerobic conditions at the sediment/water interface and possible water stratification.

A log break defines the interval top. The age is based upon the following palynological criteria:

- the appearance of *Cannosphaeropsis thula* at 2665m(SWC).
- the presence, at 2703m(SWC), of *Aldorfia spongiosa* and of *Batioladinium pomum* at 2710m.

MICROPALAEONTOLOGY

The appearance of a specimen of *Haplophragmoides infracallovienensis* at 2662.50m provides evidence for the penetration of the Kimmeridge Clay Formation. The fact that the log break marking the top of this formation is at 2664m suggests that there is a slight discrepancy between the drillers depths and log depths. Rare specimens of *Haplophragmoides* spp., exhibiting the brown preservation associated with the Kimmeridge Clay Formation, are the only other in situ foraminifera present, although caved Tertiary and Cretaceous forms are frequently seen. The presence of common specimens of the radiolarian *Lithostrobos* spp. at and below 2685m may prove of use for local correlations.

PALYNOLOGY

The palynological analysis of ditch cuttings and sidewall core samples from this interval yielded microplankton dominated assemblages, typical of the Kimmeridge Clay Formation in composition and preservation. Elements of the *Hystriosphera orbifera* group and the prasinophycean algae *Pterospermella* spp. and large leiospheres occur consistently, and are sometimes numerous to common.

A specimen of *Cannosphaeropsis thula* at 2665m(SWC) indicates an age no younger than the early Ryazanian, *Cannosphaeropsis* sp.A Subzone (VC). The subsequent occurrence of *Aldorfia spongiosa* at 2703m(SWC), a dinocyst with a stratigraphic range of early Ryazanian - late Volgian, supports this age assignment, as does the occurrence of *Batioladinium pomum* at 2710m.

VI (4) JURASSIC

INTERVAL 2740m-2770m(SWC); LATE JURASSIC, MIDDLE - EARLY VOLGIAN

Lithostratigraphic Unit:

Humber Group (part);

 Kimmeridge Clay Formation (part).

Environment: marine, inner shelf, with anaerobic/dysaerobic bottom conditions and partial water mass stratification.

The age and upper limit of the interval are defined upon the following palynological criteria:

- the appearance of *Glossodinium dimorphum* at 2740m.
- the subsequent occurrences of specimens of *Gonyaulacysta longicornis* and *Oligosphaeridium pulcherrimum* sensu Ioannides et al. in the ditch cuttings sample at 2770m.

MICROPALAEONTOLOGY

The microfaunas recovered from this interval are essentially the same as those seen above, with rare *Haplophragmoides* spp. being recorded together with specimens of *Lithostrobos* spp.

PALYNOLOGY

Three ditch cuttings samples analysed from this interval continued to yield abundant amorphous organic matter. The palynofloras recovered were dominated by microplankton; characteristic elements of the assemblages including members of the *Systematophora areolaria* and *Hystriospharina orbifera* groups, *Tubotuberella apatela*, and prasinophycean algae such as *Pterospermella* spp. and *Tasmanites* spp. Bisaccate pollen (abundant at 2762.50m), *Cerebroipollenites mesozoicus* and *Callialasporites dampieri* are typical elements of the miospore suite.

The appearance of *Glossodinium dimorphum* at 2740m indicates an age no younger than the middle Volgian, *Dichadogonyaulax pannea* Subzone (VIIA). Evidence for the presence of sediments of early Volgian age is seen only in the ditch

cuttings sample at 2770m, where specimens of *Oligosphaeridium pulcherrimum* sensu Ioannides et al. and *Gonyaulacysta longicornis* are encountered. The combined occurrence of these two taxa characterises deposits assignable to the *Egmontodinium polyplacophorum* Subzone (IXA) of the *G. longicornis* Zone (IX).

INTERVAL 2770m(SWC)-2782.50m; LATE JURASSIC, KIMMERIDGIAN

Lithostratigraphic Units:

Humber Group (part);

 Kimmeridge Clay Formation (part): 2770m(SWC)-2780m(log),

 Heather Formation (part): 2780m(log)-2782.50m.

Environment: marine, inner shelf with aerobic/dysaerobic bottom conditions.

The age and upper limit of the interval are defined upon the following palynological criterion:

- the appearance of *Gonyaulacysta jurassica* at 2770m(SWC).

MICROPALAEONTOLOGY

Rare *Haplophragmoides* spp. and specimens of *Lithostrobus* spp. continue to be recorded from this thin interval.

PALYNOLOGY

A miospore dominated palynoflora was recovered from the sidewall core at 2770m, characterised by common undifferentiated bisaccates.

The appearance of *Gonyaulacysta jurassica* indicates an age no younger than the Kimmeridgian, *Scriniodinium luridum* Subzone (IXB) of the *Gonyaulacysta longicornis* Zone (IX).

A single specimen of *Lycospore* sp. is present reflecting reworking of Carboniferous deposits.

INTERVAL 2782.50m-2855m(SWC); LATE JURASSIC, EARLY KIMMERIDGIAN - MIDDLE OXFORDIAN

Lithostratigraphic Unit:

Humber Group (part);

Heather Formation (part).

Environment: marine, inner shelf with aerobic/dysaerobic bottom conditions.

The upper limit and age of the interval are based upon the following palynological criteria:

- the appearance of *Scriniodinium crystallinum* and the occurrence of numerous specimens of *Gonyaulacysta jurassica* at 2782.50m.
- the subsequent occurrence of *Leptodinium mirabile* between 2812.50m and 2842.50m.
- the appearance of *Nannoceratopsis pellucida* at 2842.50m.

MICROPALAEONTOLOGY

The microfaunas of this interval are essentially as seen above. There is, however, a change near the base where *Lenticulina varians* appears suggesting that the bottom waters were more oxygenated in the lower part of this interval. The appearance of this species is consistent with an early Kimmeridgian - middle Oxfordian age.

PALYNOLOGY

The ditch cuttings and single sidewall core samples analysed from this interval yielded palynofloras dominated by gymnosperm pollen, including common to abundant undifferentiated bisaccates, *Cerebropollenites mesozoicus* and *Callialasporites dampieri*. Pteridophyte miospores are less numerous and less diverse. The microplankton assemblages are moderately diverse and are characterised by *Gonyaulacysta jurassica*, microforaminiferal test linings and *Sirmiodinium grossii* (which becomes numerous to common below 2812.50m).

Scriniodinium crystallinum makes its first downhole appearance at 2782.50m, indicating an age no younger than the early Kimmeridgian/late Oxfordian *Scriniodinium crystallinum* Zone (XI). The simultaneous increase in numbers of *Gonyaulacysta jurassica* is also characteristic of this age.

Nannoceratopsis pellucida is present in the deepest sample in this interval (2842.50m). This dinocyst ranges no younger than middle Oxfordian in this area of the North Sea, and indicates that the basal part of this interval is no younger than the early part of the *Compositosphaeridium costatum* Subzone (XIIB). The presence of *Leptodinium mirabile* from 2812.50m to 2842.50m (its last downhole occurrence) supports this conclusion, since this dinocyst ranges no older than the middle Oxfordian, *Acanthaulax spinosissima* Zone (XIII).

INTERVAL 2855m(SWC)-2887.50m; LATE JURASSIC, EARLY OXFORDIAN

Lithostratigraphic Unit:

Humber Group (part);

Heather Formation (part).

Environment: marine, inner shelf with aerobic bottom conditions.

The age and top of the interval are defined upon the following palynological criteria:

- the appearance of abundant specimens of *Adnatosphaeridium aemulum* and *Acanthaulax spinosissima* at 2855m(SWC).
- the subsequent occurrence of *Atopodinium prostaticum* at 2857.50m.
- the presence of *Wanaea fimbriata*, *W. thysanota* and *Crussolia deflandrei* at 2870m(SWC).

MICROPALAEONTOLOGY

Microfaunas from this interval are generally poor. The appearance of *Haplophragmoides canui* at 2880m could represent the top of the *Haplophragmoides canui* Assemblage, but a much better top is seen in the underlying interval with the influx of *Haplophragmoides* spp. and this has been preferred. The suggestion that this assemblage could be present in the base of this interval means the possibility that Callovian deposits are present from 2880m and should not be ignored.

PALYNOLOGY

Two sidewall core and two ditch cuttings samples were analysed from this interval. The sidewall core sample at 2855m yielded a very poorly preserved but diverse and numerous dinocyst palynoflora. Abundant *Adnatosphaeridium aemulum* and *Acanthaulax spinosissima* together with common *Sirmiodinium grossii* and *Gonyaulacysta jurassica* dominate.

From 2857.50m to 2870m(SWC) less diverse assemblages were recorded, characterised by numerous to common *A. aemulum* and *S. grossii* with rare to numerous *Gonyaulacysta cladophera* and consistently occurring *G. jurassica*.

Throughout the entire interval the miospore assemblages are dominated by gymnosperm pollen, particularly undifferentiated bisaccates and *Cerebropollenites mesozoicus*.

The occurrence of abundant specimens of *A. aemulum* and *A. spinosissima* at 2855m(SWC) is characteristic of an early Oxfordian age, no younger than the *Gonyaulacysta areolata* Subzone (XIVA) of the *Wanaea digitata* Zone (XIV). An early Oxfordian age is substantiated at 2857.50m by the occurrence of *Atopodinium prostaticum*.

An age equivalent to the *Wanaea fimbriata* Subzone (XIVB) is indicated at 2870m(SWC) by the occurrence of the subzonal index taxon. Within the same assemblage specimens of *Wanaea thysanota* and *Crussolia deflandrei* are encountered substantiating an age no younger than early Oxfordian. The latter taxon has, to date, only been recorded from the Tethyan Province (Southern France). Its record in this well is considered to reflect the existence of open marine connections between this area of the Viking Graben and the Tethyan Province to the south.

INTERVAL 2887.50m-2917.50m; LATE JURASSIC, LATE CALLOVIAN

Lithostratigraphic Unit:

Humber Group (part);

Heather Formation (part).

Environment: marine, inner shelf with aerobic bottom conditions.

The upper limit and age of the interval are defined upon the following palynological and micropalaeontological criteria:

- the appearance of *Mendicodinium groenlandicum* at 2887.50m.
- an increase in agglutinating foraminifera at 2890m and the subsequent occurrence of abundant *Haplophragmoides* spp., including *H. canui* from 2900m.
- the appearance of common specimens of *Pareodinia ceratophora* at 2902.50m.

MICROPALAEONTOLOGY

There is a noticeable increase in agglutinating foraminifera at 2890m followed by an influx of *Haplophragmoides* spp., including *H. canui* at 2900m. Abundant *Haplophragmoides* spp. continue to occur through the remainder of the interval. This microfauna is characteristic of the *Haplophragmoides canui* Assemblage of Callovian age.

PALYNOLOGY

Miospore dominated palynofloras were recovered from the ditch cuttings samples at 2887.50m and 2902.50m, characterised by common undifferentiated bisaccates. Relative to the overlying early Oxfordian interval, a marked change in the preservation of kerogen and reduction in both diversity and numbers of microplankton is encountered.

Specimens of *Mendicodinium groenlandicum* appear at 2887.50m, tentatively suggesting an age no younger than the late Callovian, *M. groenlandicum* Subzone (XIVC) of the *Wanaea digitata* Zone (XIV). Although this taxon possesses a total stratigraphic range of Ryazanian - Callovian, it usually achieves its numerical acme within the late Callovian. The appearance of specimens of *M. groenlandicum* at 2887.50m is, therefore, considered to probably reflect this event. Within the same assemblage, a specimen of *Pareodinia evittii* is recorded tentatively substantiating an age no younger than Callovian.

A feature of potential correlative value is the presence, at 2902.50m, of common *Pareodinia ceratophora*, a well documented intra-Callovian palynoevent.

INTERVAL 2917.50m-2970m; LATE JURASSIC, MIDDLE CALLOVIAN

Lithostratigraphic Unit:

Humber Group (part);

Heather Formation (part).

Environment: marine, inner shelf with aerobic bottom conditions.

The age and upper limit of the interval are based upon the following palynological criteria:

- the appearance of the *Lithodinia cristulata/caytonensis* plexus at 2917.50m.
- the presence of common specimens of *Chytroeisphaeridia hyalina* at 2932.50m.
- the occurrence, at 2932.50m, of *Stephanelytron scarburghense*.
- the subsequent presence of *Chytroeisphaeridia cerastes* at 2947.50m.

MICROPALAEONTOLOGY

Abundant *Haplophragmoides* spp. including *H. canui*, continue to occur in this interval with a further increase in the numbers recorded at 2950m. The *Haplophragmoides canui* Assemblage is, therefore, represented again and continues to indicate a Callovian age.

PALYNOLOGY

Three ditch cuttings samples and one sidewall core sample were analysed for palynology from this interval.

Gymnosperm pollen grains were the most common elements of the palynofloras and became increasingly abundant towards the base of the section. Microplankton assemblages recovered were moderately diverse. *Pareodinia ceratophora* is common at 2917.50m but only rare below this depth. *Chytroeisphaeridia hyalina* is common at 2932.50m becoming increasingly rare below this depth. *Gonyaulacysta jurassica*, *G. cladophora* and *Lithodinia* spp. occurred consistently throughout the interval.

The appearance of members of the *Lithodinia cristulata/caytonensis* plexus at 2917.50m indicates an age no younger than the middle Callovian, *Polystephanephorus paracalathus* Zone (XV).

The subsequent occurrence of common specimens of *Chytroeisphaeridia hyalina* at 2932.50m indicates the penetration of sediments no younger than the *Nannoceratopsis pellucida* Subzone (XVB). The last downhole occurrence of *Stephanelytron scarburghense* at the same horizon and *Chytroeisphaeridia cerastes* at 2947.50m supports this age assignment, as these dinocysts range no older than middle Callovian, *N. pellucida* Subzone (XVB) of the *Polystephanephorus paracalathus* Zone (XV).

INTERVAL 2970m-2977.50m(log); LATE JURASSIC, EARLY CALLOVIAN.

Lithostratigraphic Unit:

Humber Group;

Heather Formation (part).

Environment: marine, inner shelf with aerobic bottom conditions.

The upper limit and age of the interval are defined upon the following palynological criteria:

- the appearance of *Sentusidinium varispinosum* at 2970m.
- the occurrence of ?caved specimens of *Hystrihogonyaulax pectinigerum* at 2977.50m.

An unconformity is considered to be present at the base of this interval, resulting in the absence of deposits of late - middle Bathonian age.

MICROPALAEONTOLOGY

The single sample analysed from this interval shows a significant decline in the number of foraminifera recorded when compared to the middle Callovian. No new age diagnostic forms were seen.

PALYNOLOGY

The palynoflora recovered from the single ditch cuttings sample analysed from this interval (2970m) was dominated by abundant gymnosperm pollen, such as *Callialasporites* spp., *Cerebropollenites mesozoicus* and undifferentiated bisaccates. The microplankton assemblage is characterised by numerous specimens of *Sirmiodinium grossii*, *Sentusidinium varispinosum* and *Lithodinia* spp..

The appearance of *S. varispinosum* at 2970m indicates the penetration of early Callovian sediments, equivalent to the *Dichadogonyaulax gochtii* Subzone (XVIA) and lowest part of the *Nannoceratopsis pellucida* Subzone (XVB). The presence of numerous specimens of this taxon is normally a feature encountered within the *D. gochtii* Subzone (XVIA) and may indicate an age no younger than this subzone at 2970m.

A specimen of *Hystrihogonyaulax pectinigerum* is present at 2977.50m (at the top of the underlying interval) within a palynoflora significantly modified by caved Callovian palynomorphs. Although *Hystrihogonyaulax pectinigerum* may occur in sediments of pre-Callovian age in onshore Northwestern Europe, the good preservation of the specimen and its local restriction to sediments of early Callovian age in this area of the North Sea, strongly suggest derivation from this early Callovian interval.

INTERVAL 2977.50m(log)-2991m(log); MIDDLE JURASSIC, EARLY BATHONIAN - LATE BAJOCIAN

Lithostratigraphic Unit:

Brent Group (part):

Tarbert Formation.

Environment: marginal marine/inner shelf with strong terrestrial influences.

The upper limit is defined upon a log break. The age is based upon the following palynological criterion:

- the presence, at 2977.50m, of numerous specimens of *Dissiliodinium?* "granulatum".

Regional considerations suggest the possibility of an unconformity or stratigraphic hiatus at the base of the interval resulting in the absence of some deposits of late Bajocian and middle Bajocian ages.

MICROPALAEONTOLOGY

This interval is barren of in situ foraminifera.

PALYNOLOGY

A single ditch cuttings sample and two core samples were analysed for palynology from this interval. The palynoflora recovered from the ditch cuttings sample is strongly modified by caving from the overlying Callovian interval. The appearance of numerous specimens of *Dissiliodinium?* "granulatum" at 2977.50m is characteristic of the early Bathonian - late Bajocian in this area of the North Sea. A specimen of *Hystrichogonyaulax pectinigerum* present at the same horizon is considered to be derived, via caving, from the overlying early Callovian interval.

The core sample at 2980.70m yielded a palynoflora dominated by the abundant occurrence of indeterminate algal sacs. Rare gymnosperm pollen were also recorded.

The core sample at 2981.33m produced a more numerous palynoflora with abundant bisaccates and Perinopollenites elatodes present. The only examples of marine microplankton in the core pieces were a faintly granular specimen of Sentusidinium sp. and Caddasphaera halosa.

INTERVAL 2991m(log)-3070m; MIDDLE JURASSIC, MIDDLE - EARLY BAJOCIAN

Lithostratigraphic Units:

Brent Group (part);

Ness Formation: 2991m(log)-3034m(log).

Unassigned Unit: 3034m(log)-3070m.

Environment: initially marginal marine/inner shelf, becoming progressively marginal marine/deltaic with strong terrestrial influences towards the top.

The upper limit is defined upon a log break. The age is based upon the following palynological criteria:

- the presence of common specimens of *Leiofusa "solida"* at 2986.63m (uncorrected core depth).
- the simultaneous appearance of numerous specimens of *Nannoceratopsis gracilis* and *N. senex*, in association with *Chasmatosporites* spp.

MICROPALAEONTOLOGY

No in situ foraminifera occur in this interval reflecting the marginal marine environment indicated by palynological evidence.

PALYNOLOGY

Core pieces, sidewall cores and ditch cuttings samples were analysed from this interval. Gymnosperm pollen formed the dominant element of the palynofloras recovered, with numerous to abundant undifferentiated bisaccates, *Callialasporites dampieri*, *Perinopollenites elatoides* and *Cerebropollenites mesozoicus* occurring consistently. With the exception of the core sample at 2986.63m (uncorrected depth) microplankton assemblages are sparse, with sporadic occurrences of *Nannoceratopsis* sp. *Tasmanites* spp. and the freshwater alga *Botryococcus* spp.

At 2986.63m(core, uncorrected depth) the microplankton assemblage is essentially composed of three forms: the distinctive netromorph acritarch *Leiofusa "solida"*, and the dinocysts *Nannoceratopsis gracilis* and *N. senex*. *Leiofusa "solida"* usually marks the penetration of sediments equivalent to the upper part of the Ness Formation in this region of the Oseberg Field complex. The simultaneous appearance of numerous specimens of *N. senex* indicates an age no

younger than middle Bajocian, as does the first appearance of Chasmatosporites spp. at this same depth: a miospore that is considered to range no younger than Bajocian.

The increased incidence of inertinitic and vitrinitic kerogen within this interval reflects the proximity of terrestrial, swamp-type macrofloras, an interpretation supported by the increased diversity of pteridophyte miospores recorded.

An impoverished palynoflora was recovered from the sidewall core at 3050.70m preventing accurate dating of the lower part of this interval. In addition, ditch cuttings samples yielded assemblages significantly modified by caved taxa from the Callovian intervals and the upper part of this interval (Ness Formation).

INTERVAL 3070m-3186m(SWC); MIDDLE - EARLY JURASSIC, EARLIEST BAJOCIAN - MIDDLE TOARCIAN

Lithostratigraphic Units:

Unassigned Unit: 3070m-3116m(log).

Dunlin Group (part);

Drake Formation: 3116m(log)-3186m(SWC).

Environment: initially marine, inner shelf with aerobic/dysaerobic bottom conditions, becoming progressively marginal marine/inner shelf.

The age and upper limit of the interval are defined upon the following palynological criteria:

- the occurrence of *Mancodinium semitabulatum* at 3070m.
- the presence of a questionable specimen of *Scriniocassis weberi*, at 3080m.
- the association of *S. weberi*, *Nannoceratopsis ambonis* and *Comparodinium punctatum* at 3117m(SWC).

MICROPALAEONTOLOGY

The microfaunas of this interval are extremely impoverished and it is suspected that the rare specimens of agglutinated foraminifera recorded are caved from the Late Jurassic. The occurrence of a single specimen of *Trochammina* sp. at 3120m may reflect the presence of late - middle Toarcian deposits at this depth.

PALYNOLOGY

Palynofloras recovered from the ditch cuttings samples within this interval are significantly modified by caved taxa from the overlying Late Jurassic, particularly Callovian, intervals, resulting in the recovery of impoverished in situ assemblages. In addition, apart from the sidewall core at 3117m, palynofloras recovered from the sidewall cores are also impoverished. As a result, the dating of this interval is based upon meagre evidence, hence the wide ranging age assigned.

Assemblages are dominated by miospores throughout, characterised by the common to abundant occurrences of undifferentiated bisaccates, *Perinopollenites elatoides*, the *Baculatisporites/Osmundacidites* group, *Cerebropollenites mesozoicus*, *Deltoidospora* spp., *Callialasporites* spp. and *Chasmatosporites* spp.

Common specimens of *Nannoceratopsis senex* are encountered at and below 3117m(SWC), becoming abundant at 3172.50m.

A single specimen of *Mancodinium semitabulatum* was recorded at 3070m and based upon regional considerations is considered to indicate an age no younger than the lower part of the early Bajocian or even a Toarcian age. This distinctive taxon, which can extend up into deposits of middle Bajocian age, normally, in the Oseberg Field complex, does not occur in sediments of post-earliest Bajocian age. This age assignment is tentatively supported by the presence of a questionable specimen of *Scriniocassis weberi* at 3080m, a species which does not extend into post-earliest Bajocian deposits.

The sidewall core at 3117m yielded a relatively rich assemblage containing the association of *Scriniocassis weberi*, *Nannoceratopsis ambonis* and *Comparodinium punctatum*, which on local correlation suggests a late - middle Toarcian age.

An age no younger than the "Sphaeromorphs" Subzone (XVIIIC) of the *Nannoceratopsis gracilis* Zone (XVIII) is indicated at 3172.50m by the appearance of numerous clusters of "sphaeromorph" acritarchs.

Of additional interest is the occurrence, at 3117m(SWC), of recycled miospores including striate bisaccates (Permo-Triassic) and *Striatoabieites aytugii* (Carnian - Scythian). Their occurrence at this horizon may be of local correlative value, reflecting a phase of introduction of reworked material from the source area.

INTERVAL 3186m(SWC)-3220m(SWC); EARLY JURASSIC, EARLY TOARCIAN

Lithostratigraphic Unit:

Dunlin Group (part);

Drake Formation (part).

Environment: marine, inner shelf with anaerobic/dysaerobic bottom conditions.

The upper limit and age of this interval are based upon the following palynofloral criterion:

- the occurrence of abundant specimens of "sphaeromorph" acritarchs at 3186m(SWC).

MICROPALAEONTOLOGY

Impoverished microfaunas, as seen in the overlying interval, continue to be recorded.

PALYNOLOGY

One ditch cuttings and one sidewall core sample were analysed from this interval. Although gymnosperm pollen continue to be the dominant element of the palynofloras, a significant change in their composition was recorded with an influx of common specimens of Chasmatosporites spp. with Perinopollenites elatoides and Callialasporites dampieri becoming very rare.

The microplankton assemblage in the sidewall core at 3186m is characterised by the occurrence of abundant "sphaeromorph" acritarch clusters. This palynoevent is indicative of the penetration of sediments equivalent to the early Toarcian part of the "Sphaeromorphs" Subzone (XVIIIIC).

INTERVAL 3220m(SWC)-3418m; EARLY JURASSIC, LATE PLIENSBAKIAN - SINEMURIAN

Lithostratigraphic Units:

Dunlin Group (part): 3220m(SWC)-3387m(log);

Drake Formation: 3220m(SWC)-3241m(log),

"Sand Unit": 3241m(log)-3288m(log),

"Shale Unit": 3288m(log)-3387m(log).

Statfjord Formation: 3387m(log)-3418m.

Environment: initially deltaic/marginal marine, becoming inner shelf with strong terrestrial influences.

The upper limit and age of this interval are defined upon the following palynological criteria:

- the occurrence of a latest Pliensbachian palynoflora at 3220m(SWC).
- the subsequent appearance of *Cerebropollenites* cf. *thiergartii* at 3232.50m.

MICROPALAEONTOLOGY

This interval is barren of in situ microfauna, probably as a result of its marginal marine environmental setting.

PALYNOLOGY

From 3220m(SWC) to 3265m (ditch cuttings) the palynofloras examined were dominated by gymnosperm pollen, particularly undifferentiated bisaccates, *Chasmatosporites* spp. and *Cerebropollenites* spp. Marine microplankton form a prominent part of the palynoflora in the sidewall core sample at 3220m. This is reduced to very rare acanthomorph acritarchs and prasinophycean algae in the ditch cuttings sample from 3265m, whilst in the sidewall core sample at the same depth an exclusively terrestrial derived assemblage was encountered. A single specimen of *Tasmanites* sp. was present at 3297.50m(SWC) indicating weak/fluctuating salinities.

No in situ marine microplankton were recorded from 3307.50m to 3418m. Within this interval pteridophyte miospores became more diverse and numerous, with *Deltoidospora* spp. and elements of the *Baculatisporites*/*Osmundacidites* group frequently numerous to common and *Klukisporites* spp., *Verrucosisporites* spp. and *Dictyophyllidites* spp. occurring consistently. Gymnosperm pollen still form the

dominant part of the palynofloras, however, characterised by numerous to common undifferentiated bisaccates and *Cerebropollenites* cf. *thiergartii* with consistently occurring *Chasmatosporites* spp.

At 3220m(SWC) the outgoing of common to abundant "sphaeromorph" acritarchs indicates a late Pliensbachian age. The base of the range of *Nannoceratopsis tricerata* at the same depth indicates an age equivalent to the upper part of the *Luehndea spinosa* Subzone (XVIIID), an age assignment supported by the last downhole occurrence of *N. senex* and *N. gracilis*, neither of which range any older than late Pliensbachian. Further evidence for a late Pliensbachian age comes from the appearance of *Cerebropollenites* cf. *thiergartii* at 3232.50m and its common occurrence at 3265m(SWC), a feature characteristic of sediments of this age in this part of the North Sea.

Although no positive biostratigraphic evidence for a Sinemurian age was recovered from this interval the penetration of sands assigned to the Statfjord Formation suggests that, on regional grounds, the interval may be as old as Sinemurian in the lower part.

Rare occurrences of striate bisaccates (3297.50m(SWC) and 3376m(SWC)) indicate the recycling of Permo-Triassic sediments within this interval.

GEOLOGICAL HISTORY

SINEMURIAN TO TOARCIAN

The oldest sediments encountered are fine to coarse grained sandstones deposited within a fluvio-deltaic setting, with the possibility of minor marine influences. Sand deposition ceased abruptly, probably within the late Sinemurian, resulting in the accumulation of shales within a marginal marine to inner shelf setting. Within the Pliensbachian, a significant influx of arenaceous detritus occurred within a proximal, marine location. During the late Pliensbachian, shale deposition resumed and continued into the early Toarcian, during which time anoxic bottom conditions prevailed.

TOARCIAN TO LATE BAJOCIAN-
EARLY BATHONIAN

Within the late - middle Toarcian period a major influx of sand occurred which continued into the early Bajocian. These deposits are considered to probably have been laid down within an inner shelf/marginal marine environment with strong terrestrial influence. Continued regression during the early Bajocian resulted in the progradation of the delta-top facies of the Ness Formation. This style of sedimentation continued into the middle Bajocian.

A late Bajocian transgressive phase subsequently inundated the delta-top, following a brief phase of non-deposition and/or erosion, and led to the accumulation of marine sands. This style of

	sedimentation continued to the early Bathonian.
MIDDLE - LATE BATHONIAN	ABSENT - due to either deposition with subsequent erosion or non-deposition.
CALLOVIAN TO LATE VOLGIAN- EARLY RYAZANIAN	Sedimentation resumed in response to the widespread early Callovian transgressive phase with the accumulation of shales and limestones in an inner shelf environment with aerobic bottom conditions. Deposition appears to have been continuous into the early Ryazanian with the accumulation of shales under progressively anaerobic/dysaerobic bottom conditions with the establishment of a stratified water column.
LATE RYAZANIAN TO HAUTERIVIAN	ABSENT - due to non-deposition or deposition with subsequent erosion.
BARREMIAN TO MIDDLE ALBIAN	Sedimentation resumed within the Barremian. Deposition was probably intermittent with very slow subsidence rates throughout the Aptian to middle Albian, within a shallow, oxygenated, inner shelf location.
LATE ALBIAN TO CONIACIAN	ABSENT - due to non-deposition or deposition with subsequent erosion.
SANTONIAN TO MAASTRICHTIAN- LATE MAASTRICHTIAN	Deposition resumed during the Late Cretaceous resulting in the accumulation of a predominantly argillaceous sequence within an inner to outer shelf location.

PALAEOCENE

Following a minor unconformity, probably due to the effects of North Atlantic rifting, sedimentation resumed within the Early Palaeocene, Danian. Deposition was continuous into the Late Palaeocene under outer shelf to upper bathyal depths. Towards the end of the Late Palaeocene a significant volcanic event, associated with further North Atlantic rifting, resulted in the deposition of tuffaceous shales.

SELECTED REFERENCES

- BARTENSTEIN, H. et al. 1962. Leitfossilien der Mikropaleontologie. Gebrüder Borntraeger, Berlin.
- BATE, R. & ROBINSON, E. 1978. A stratigraphical index of British Ostracoda. British Micropalaeontological Series.
- BERGGREN, W.A. 1971. Tertiary boundaries and correlations. In: FUNNELL, B.M. & RIEDEL, W.R. (Eds.), Micropalaeontology of the Oceans. Camb. Univ. Press.
- BLOW, W.H. 1979. The Cainozoic Globigerinida. E.J. Brill. Leiden.
- DAVEY, R.J. 1979. The Stratigraphic distribution of Dinocysts in the Portlandian (Latest Jurassic) to Barremian (Early Cretaceous) of Northwest Europe. AASP. Cont. Series No. 5B.
- DAVEY, R.J. 1982. Dinocyst stratigraphy of the latest Jurassic to Early Cretaceous of the Haldager No. 1 borehole, Denmark. Danm. geol. Unders., Series B, No. 6.
- DEEGAN, C.E. & SCULL, B.J. (Compilers). 1977. A proposed standard lithostratigraphic nomenclature for the Central and Northern North Sea. Rep. Inst. Geol. Sci., No. 77/25; Bull. Norw. Petrol. Direct., No. 1.
- FINSTAD, K.G. & SELLEY, R.C. (Eds.). 1975. Proceedings, Jurassic Northern North Sea Symposium. Norwegian Petroleum Society.
- FINSTAD, K.G. & SELLEY, R.C. (Eds.). 1977. Proceedings, Mesozoic Northern North Sea Symposium. Norwegian Petroleum Society.
- GRADSTEIN, F.M. & BERGGREN, W.A. 1981. Flysch-Type Agglutinating Foraminifera and the Maestrichtian to Paleogene History of the Labrador and North Seas. Marine Micropaleontology, Vol. 6.

ILLING, L.V. & HOBSON, G.D. 1980. Petroleum Geology of the Continental Shelf of North-West Europe. Institute of Petroleum.

IOANNIDES, N.S., STAVRINOS, G. & DOWNIE, C. (1976) 1977. Kimmeridgian microplankton from Clavell's Hard, Dorset, England. Micropaleontology, Vol. 22.

JENKINS, D.G. & MURRAY, J.W. (Eds.). 1981. Stratigraphical Atlas of fossil foraminifera. British Micropalaeontological Society Series.

LORD, A.R. 1982. A Stratigraphical Index of Calcareous Nannofossils. British Micropalaeontological Society Series.

THUSU, B. (Ed.). 1978. Distribution of biostratigraphically diagnostic dinoflagellate cysts and miospores from the Northwest European Continental Shelf and adjacent areas. Cont. Shelf. Inst. Publ. No. 100.

TYSON, R.V., WILSON, R.C.L. & DOWNIE, C. 1979. A stratified water column environmental model for the type Kimmeridge Clay. Nature, Vol. 277.

WOODLAND, A.W. (Ed.). 1975. Petroleum and the Continental Shelf of North West Europe, Vol. 1. Geology. Institute of Petroleum.

ROBERTSON RESEARCH INTERNATIONAL LIMITED. 1979. Southern Offshore Norway: The Stratigraphy and Petroleum Geochemistry of the Jurassic to Tertiary Sediments.

ROBERTSON RESEARCH INTERNATIONAL LIMITED/ERC ENERGY RESOURCE CONSULTANTS LIMITED. 1980. The Brent Sand in the Northern Viking Graben, North Sea: A Sedimentological and Reservoir Engineering Study.

ROBERTSON RESEARCH INTERNATIONAL LIMITED. 1982. Southern Offshore Norway Phase Two Study: The Stratigraphy and Petroleum Geochemistry of the Jurassic to Tertiary Sediments.

APPENDIX 1

LITHOLOGICAL DESCRIPTIONS OF SIDEWALL CORES

Depth (m)	Recovery (mm)	Quality	Type of Analysis	Lithology
2002	frags	fair	M	<u>SANDY CLAYSTONE</u> : moderately hard, medium dark grey, highly argillaceous with traces of mica and sands, weakly calcareous and ?tuffaceous.
2006	20	fair		<u>SHALE</u> : moderately hard, weakly fissile, finely laminated, medium dark grey.
2008	17	fair		<u>SHALE</u> : moderately hard, slightly waxy, weakly fissile, medium dark grey and slightly micromicaceous.
2023	20	fair	M	<u>CLAYSTONE</u> : moderately hard, slightly waxy, finely laminated and medium dark grey.
2165	frags	fair	M	<u>CLAYSTONE</u> : moderately hard, slightly waxy, olive black to greenish black, micromicaceous.
2225	10	fair	M	<u>CLAYSTONE</u> : moderately hard, slightly waxy, medium grey to olive grey, micromicaceous and strongly calcareous.
2432	15	fair	M	<u>CALCAREOUS CLAYSTONE</u> : moderately hard, slightly waxy, olive black, micromicaceous and moderately calcareous.
2580	10	fair	MP	<u>CLAYSTONE</u> : as 2432m, but weakly calcareous.
2640	10	fair	MP	<u>CLAYSTONE</u> : moderately hard, slightly waxy, olive black to brownish black and micromicaceous.

2665	frags	fair	MP	<u>CALCAREOUS CLAYSTONE/ARGILLACEOUS LIMESTONE</u> : hard, olive black, highly argillaceous and micromicaceous.
2703	frags	fair	MP	<u>SHALE</u> : hard, subfissile, olive black to brownish black, micaceous and very weakly to non-calcareous.
2727	frags	fair	P	<u>SHALE</u> : moderately hard, earthy, brownish black, micromicaceous, silty and weakly calcareous.
2770	10	fair	P	<u>SHALE</u> : moderately hard, fissile, brownish black, micromicaceous.
2815	frags	fair	P	<u>SHALE</u> : moderately hard, fissile, olive black to dark grey, micaceous.
2855	frags	poor	P	<u>LIMESTONE</u> : hard, coarsely crystalline, brownish grey and micaceous.
2870	frags	fair	P	<u>SILTY SHALE</u> : moderately soft, subfissile, dusky yellowish brown, silty with very fine sand, micromicaceous and weakly calcareous.
2955	frags	fair	P	<u>SILTY SHALE</u> : as 2870m, but non-calcareous.
2976.30	frags	poor		sample predominantly drilling mud with fragments of <u>SHALE</u> : brownish black with <u>COAL</u> : black.
2997.50	frags	poor	P	<u>SHALE</u> : waxy, subfissile, moderately hard, brownish grey.
3025	frags	poor	P	<u>SHALE</u> : as 2997.50m with carbonaceous fragments.
3050.70	12	fair	P	<u>SANDSTONE</u> : fairly friable, pinkish grey, very fine grained, micaceous and weakly calcareous.
3085	18	fair	P	<u>SANDSTONE</u> : fairly friable, white to pinkish grey, fine to medium grained, micaceous, ?kaolinitic and weakly calcareous.

3117	frags	fair	P	<u>SANDSTONE</u> : friable, light grey, very fine grained, slightly argillaceous, micaceous, finely laminated with <u>SHALE</u> : brownish grey and highly micaceous with carbonaceous fragments, and <u>ARGILLACEOUS SANDSTONE</u> : light brownish grey to medium light grey, very fine grained and micaceous.
3158.50	frags	fair	P	<u>SANDSTONE</u> : friable, light grey, very fine to medium grained, ?kaolinitic.
3186	frags	fair	P	<u>SHALE</u> : moderately hard, weakly fissile, olive black and micromicaceous.
3220	frags	fair	P	<u>SHALE</u> : moderately hard, subfissile, brownish grey to dusky yellowish brown, slightly silty, micaceous.
3265	frags	fair	P	<u>ARGILLACEOUS SANDSTONE</u> : fairly friable, very pale orange to dark yellowish brown, very fine grained, argillaceous and micaceous, laminated with <u>SHALE</u> : dusky yellowish brown and micaceous.
3297.50	10	poor	P	<u>CLAYSTONE</u> : moderately hard, olive black, micaceous and very weakly calcareous.
3346	frags	fair	P	<u>SHALE</u> : moderately hard, subfissile, dusky yellowish brown, micaceous, locally sandy grading to <u>ARGILLACEOUS SANDSTONE</u> .
3376	frags	fair	P	<u>SHALE</u> : with minor <u>ARGILLACEOUS SANDSTONE</u> : as 3346m.
3391	18	fair	P	<u>SANDSTONE</u> : fairly friable, pinkish grey, fine to medium grained, micaceous with rare carbonaceous fragments.
3406	15	fair	P	<u>SANDSTONE</u> : as 3391m but very fine to fine grained.

3416 frags fair P SANDSTONE: friable, very light grey,
fine to coarse grained, micaceous and
slightly argillaceous, ?kaolinitic.

Key

P = Palynology
M = Micropalaeontology
frags = fragments

APPENDIX 2

LITHOLOGICAL DESCRIPTIONS OF CORE PIECES

Depth (m)	Core No.	Type of Analysis	Lithology
2980.70	1	P	<u>ARGILLACEOUS SANDSTONE</u> : well cemented, medium light grey, very fine grained, argillaceous and micaceous with carbonaceous fragments.
2981.33	1	P	<u>MICACEOUS SHALE</u> : moderately hard, fissile, medium dark grey to olive black and highly micaceous.
2986.63	1	P	<u>SHALE</u> : moderately hard, subfissile, medium dark grey and micromicaceous.
2991.85	1	P	<u>CLAYSTONE</u> : moderately hard, olive black to dark grey, micromicaceous with carbonaceous fragments.

Key

P : Palynology

GEOLOGICAL AGE		STANDARD AMMONITE ZONES	DINOCYST ZONATION			
			ZONES	SUBZONES		
LATE JURASSIC	Volgian	late	<i>lamplughii</i>	"Imbatodinium" villosum VI	<i>Egmontodinium</i> sp. A VIA	
			<i>preplicomphalus</i>		<i>Kleithrasphaeridium</i> sp. A VIB	
			<i>primitivus</i>		<i>Systematophora</i> spp. VIC	
		middle	late	<i>oppressus</i>	<i>Muderongia</i> sp. A VII	<i>Dichadogonyaulax pannea</i> VIIIA
				<i>giganteus</i>		<i>Dichadogonyaulax culmula</i> VIIIB
				<i>gorei</i>		<i>Glossodinium dimorphum</i> VIIC
			middle	<i>albani</i>	<i>Pareodinia mutabilis</i> VIII	<i>Gonyaulacysta pennata</i> VIIIA
				<i>Epipalliceras</i> sp.		<i>Gonyaulacysta jurassica</i> VIIIB
				<i>rotunda</i>		
				<i>pallasioides</i>	<i>Gonyaulacysta longicornis</i> IX	<i>Egmontodinium polyplacophorum</i> IXA
				<i>Pavlovia</i> sp.		<i>Scriniodinium luridum</i> IXB
				<i>pectinatus</i>		
	early	<i>hudlestoni</i>	<i>Gonyaulacysta cladophora</i> X			
		<i>wheatleyensis</i>				
		<i>scitulus</i>				
		<i>elegans</i>				
		<i>autissiodorensis</i>				
		<i>audoxus</i>				
	Kimmeridgian	late	<i>mutabilis</i>	<i>Scriniodinium crystallinum</i> XI	<i>Leptodinium egemeni</i> XIA	
			<i>cymodoce</i>		<i>Stephanelytron redcliffense</i> XIB	
			<i>baylei</i>		<i>Scriniodinium oxfordianum</i> XIIA	
		middle	<i>transversarium</i>	<i>Scriniodinium galeritum</i> XII	<i>Compositosphaeridium costatum</i> XIIIB	
			<i>plicatilis</i>		<i>Acanthaulax spinosissima</i> XIII	
			<i>cordatum</i>			
early	<i>marlae</i>	<i>Wanaea digitata</i> XIV	<i>Gonyaulacysta areolata</i> XIVA			
	<i>lamberti</i>		<i>Wanaea fimbriata</i> XIVB			
	<i>athleta</i>		<i>Mendicodinium groenlandicum</i> XIVC			
Callovian	late	<i>coronatum</i>	<i>Polystephanephorus paracalathus</i> XV	<i>Kalyptea stegasta</i> XVA		
		<i>jason</i>		<i>Nannoceratopsis pellucida</i> XVB		
		<i>calloviense</i>				
	middle	<i>macrocephalus</i>	<i>Pareodinia ceratophora</i> XVI	<i>Dichadogonyaulax gochti</i> XVIIA		
		<i>discus</i>				
		<i>aspidoidea</i>				
early	<i>retrocostatum</i>	<i>Nannoceratopsis spiculata</i> XVII		<i>Wanaea acollaris</i> XVIIIB		
	<i>morrisi</i>					
	<i>subcontractus</i>			<i>Gonyaulacysta filipicata</i> XVIIIC		
Bajocian	late	<i>progracilis</i>				
		<i>zigzag</i>				
		<i>parkinsoni</i>				
	middle	<i>garantiana</i>	<i>Nannoceratopsis gracilis</i> XVIII	<i>Polysphaeridium deflandrei</i> XVIIIIA		
		<i>subfurcatum</i>		<i>Mancodinium semitabulatum</i> XVIIIIB		
		<i>humphriesianum</i>		"Sphaeromorphs" XVIIIIC		
early	<i>sauzii</i>	<i>Polysphaeridium langii</i> XIX		<i>Luehndea spinosa</i> XVIIIID		
	<i>laeviuscula</i>			Unnamed subzone XIXA		
	<i>discites</i>			<i>Liasidium variabile</i> XIXB		
Toarcian	late	<i>concaevum</i>	<i>Pareodinia mutabilis</i> VIII	Unnamed subzone XIXC		
		<i>murchisonae</i>				
		<i>opalinum</i>		<i>Dapcodinium priscum</i> XIXD		
	middle	<i>levesquei</i>	<i>Rhaetogonyaulax haetica</i> XX			
		<i>thouersense</i>				
		<i>variabilis</i>				
Pliensbachian	late	<i>bifrons</i>	<i>Rhaetogonyaulax haetica</i> XX			
		<i>falciferum</i>				
		<i>tenuicostatum</i>				
	early	<i>spinatum</i>				
		<i>margaritatus</i>				
		<i>davoeri</i>				
Sinemurian	late	<i>ibex</i>	<i>Rhaetogonyaulax haetica</i> XX			
		<i>jamesoni</i>				
		<i>raricostatum</i>				
	early	<i>oxynotum</i>				
		<i>obtusum</i>				
		<i>turneri</i>				
Hettangian	<i>semicostatum</i>					
	<i>bucklendi</i>					
	<i>angulata</i>					
Rhaetian	<i>liasicus</i>					
	<i>planorbis</i>					

LITHOLOGIES

Clay	
Shale/mudstone	
Siltstone	
Sand/sandstone, very fine to medium grained	
Sand/sandstone, coarse grained to granules	
Conglomerate (with sand matrix)	
Conglomerate (without sand matrix)	
Coal/lignite	
Breccia	
Limestone (undifferentiated)	
Chalky limestone	

Dolomitic limestone	
Calcareous dolomite	
Dolomite	
Chert	
Anhydrite/gypsum	
Salt (halite)	
Potassium salts	
Concretions/nodules	
Igneous rocks, undifferentiated	
Basement, undifferentiated	
Granite	

Qualifiers

Argillaceous	
Silty/sandy	
Pebbly	
Carbonaceous	
Calcareous	
Dolomitic	
Red sediments	

Accessories

Calcite	C
Ironstone/ferruginous deposits	Fe
Glaucinite	Gl
Kaolinite	K
Phosphate	Ph
Pyrite	Py
Siderite/sphaerosiderite	S
Silica	Si

GRAIN TYPES

Oolith	
Fossils in general	
Bioclastic debris	
Mudflakes	

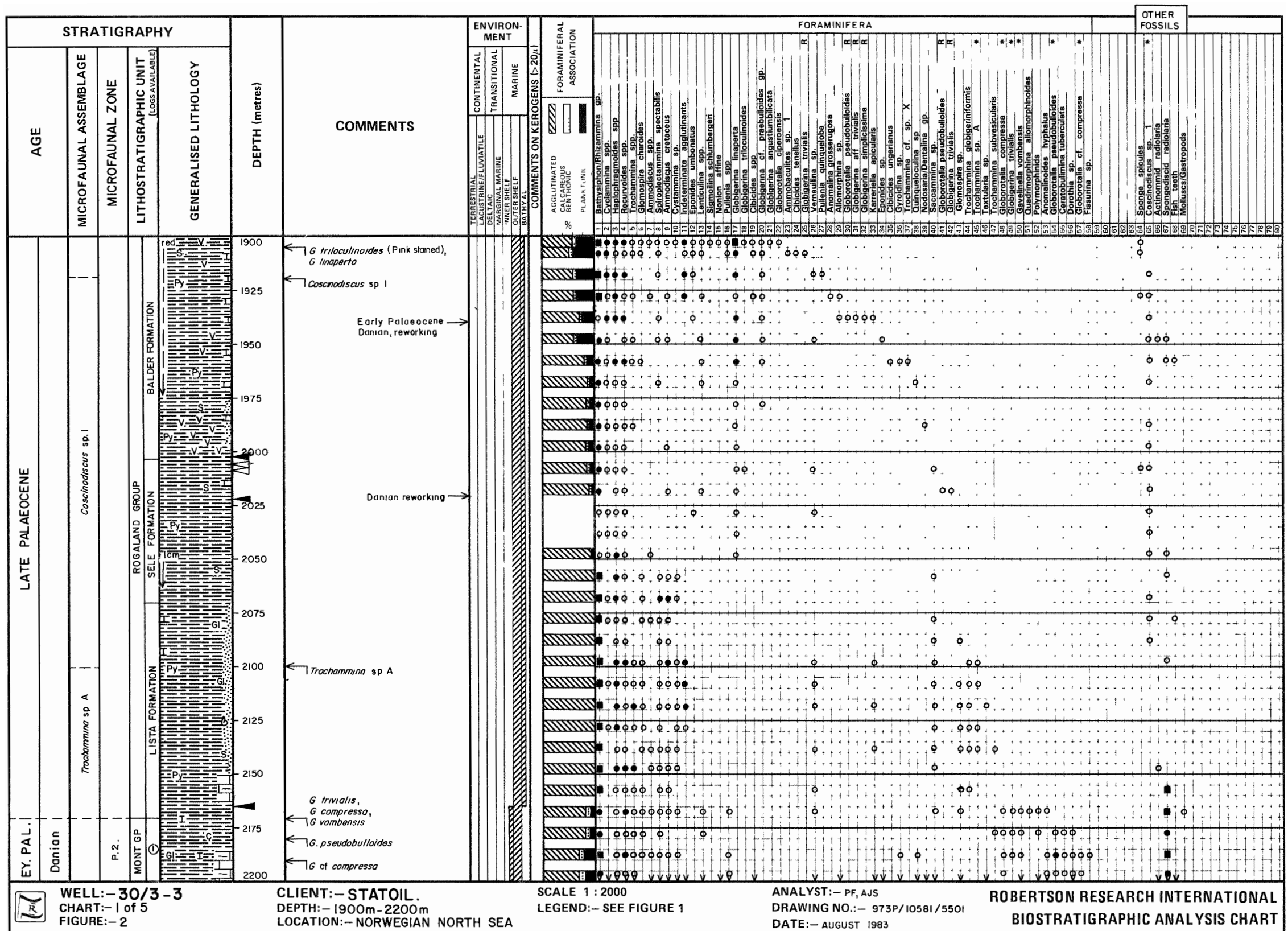
BIOSTRATIGRAPHIC SYMBOLS

Fossil Abundances	}	Present	
		Common	
		Abundant	
Diagnostic forms	*		
Caved forms	C		
Reworked forms	R		
Incoming of			
Outgoing of			
Unconformity/stratigraphic hiatus			
Faulted boundary	F—F		
LateLT., lt.		
MiddleM., m.		
EarlyEY., ey.		

Other symbols

Sample gap		
Lost circulation material	lcm	
Cement	cmt	
Turbo drilling or diamond bit drilling (Samples unsuitable for good stratigraphic analysis)	.tu	
Casing point		
Core		
Sidewall core		
Sidewall core (analysed for biostratigraphy)		
Sidewall core (no recovery)		

FIGURE 1 – Legend (edited from Robertson Research Standard Legend).



⊙ MAUREN FORMATION EQUIVALENT

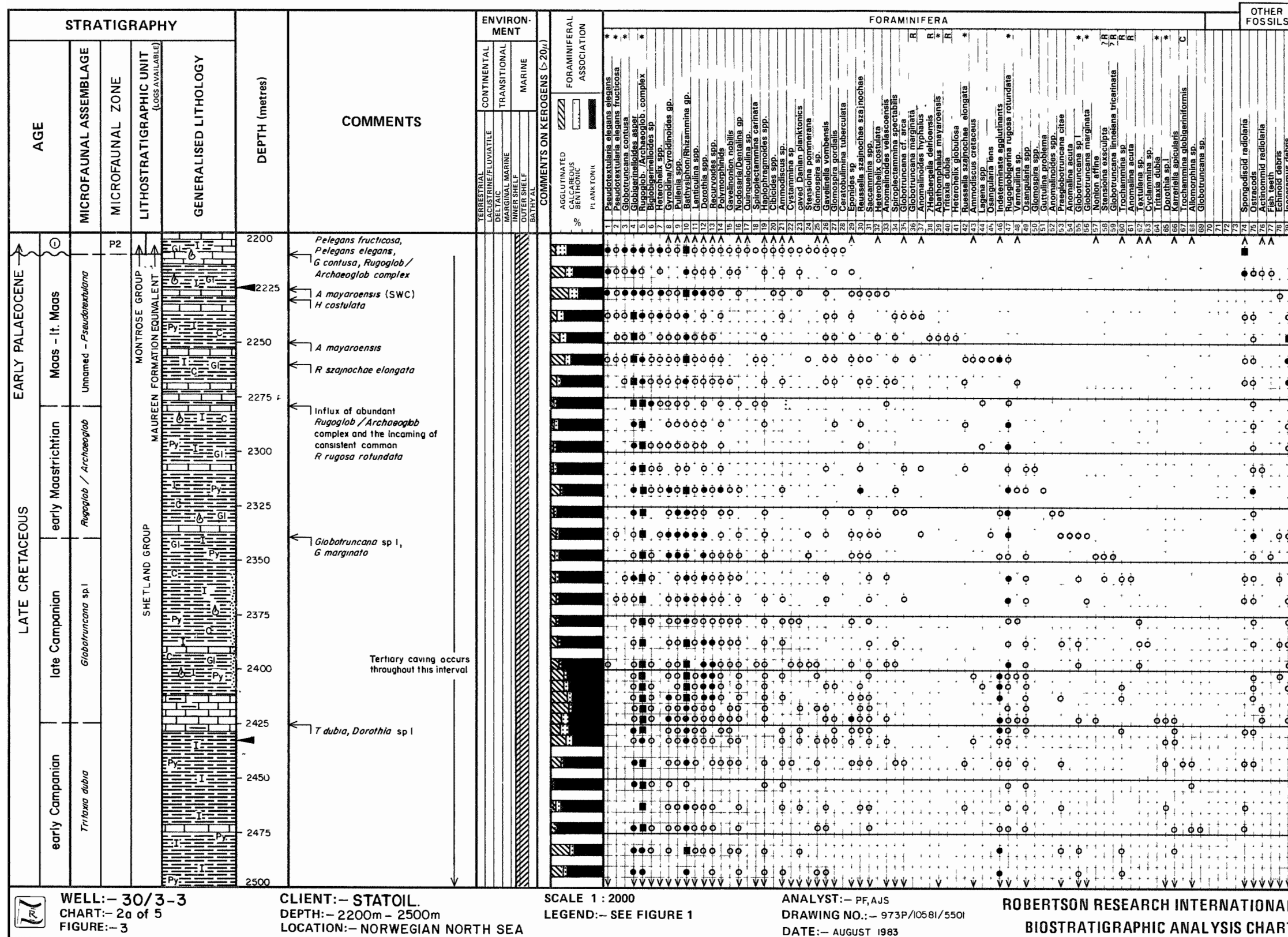
WELL: - 30/3 - 3
CHART: - 1 of 5
FIGURE: - 2

CLIENT: - STATOIL.
DEPTH: - 1900m-2200m
LOCATION: - NORWEGIAN NORTH SEA

SCALE 1 : 2000
LEGEND: - SEE FIGURE 1

ANALYST: - PF, AJS
DRAWING NO.: - 973P/10581/5501
DATE: - AUGUST 1983

ROBERTSON RESEARCH INTERNATIONAL
BIOSTRATIGRAPHIC ANALYSIS CHART



① DANIAN

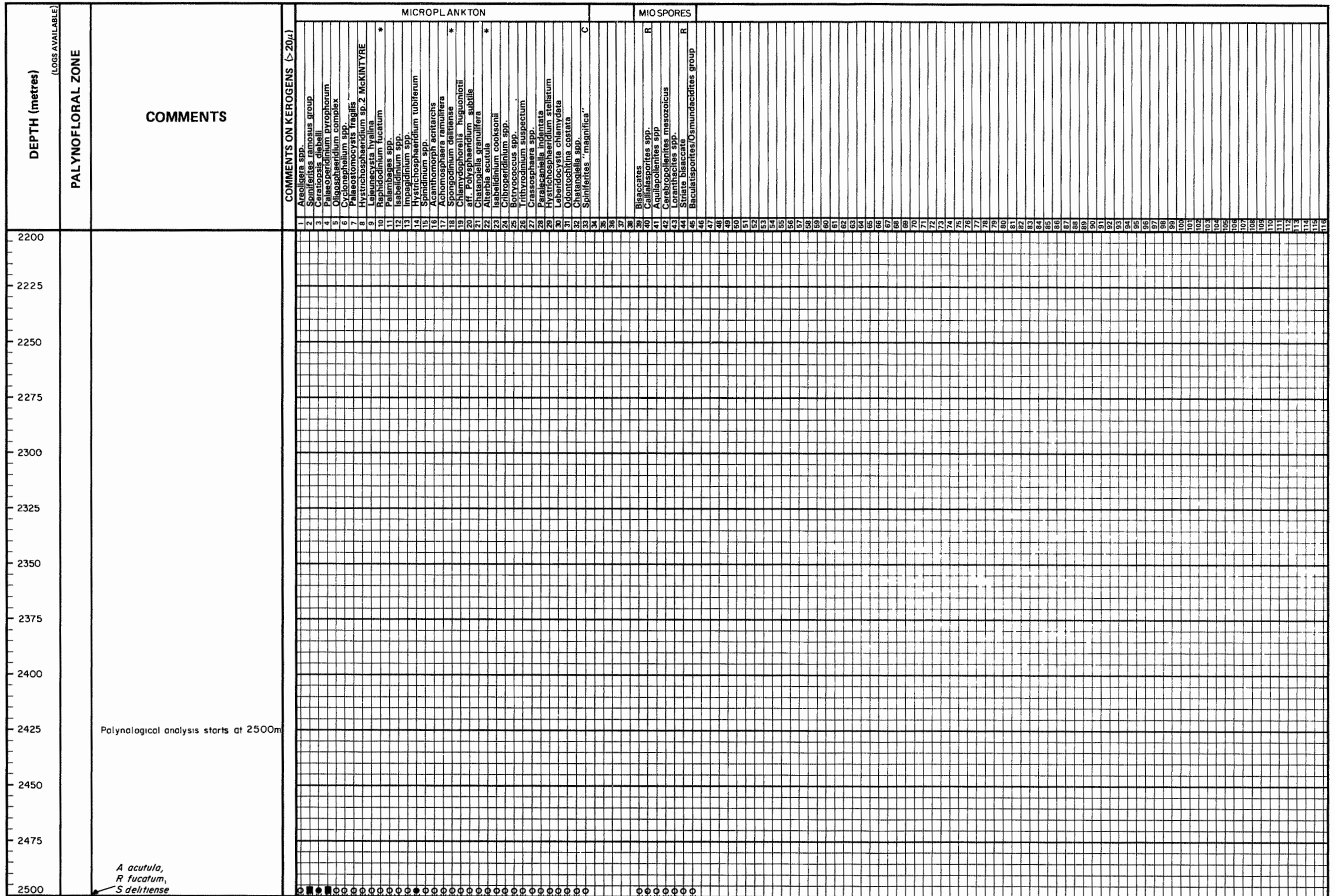
WELL:- 30/3-3
 CHART:- 2a of 5
 FIGURE:- 3

CLIENT:- STATOIL
 DEPTH:- 2200m - 2500m
 LOCATION:- NORWEGIAN NORTH SEA

SCALE 1 : 2000
 LEGEND:- SEE FIGURE 1

ANALYST:- PF,AJS
 DRAWING NO.:- 973P/10581/5501
 DATE:- AUGUST 1983

ROBERTSON RESEARCH INTERNATIONAL
 BIOSTRATIGRAPHIC ANALYSIS CHART



DEPTH (metres)
(LOGS AVAILABLE)

PALYNOFLORAL ZONE

COMMENTS

COMMENTS ON KEROGENS (>20%)		MICROPLANKTON	MIO SPORES
1	Acetabularia spp.		
2	Stauroneis group		
3	Stauroneis dihelii		
4	Pillaeoperidium pyrrophorum		
5	Oligosphaeridium complex		
6	Cyclonephalium spp.		
7	Palaestomocystis fragilis		
8	Hystriochloa sp. 2 MCKINTYRE		
9	Alpinocystis hyalina		
10	Palaeosphaera fucatum		
11	Palaeosphaera spp.		
12	Isabelidium spp.		
13	Impagidinium spp.		
14	Hystriochloa tubiferum		
15	Spinidinium spp.		
16	Acanthomorph acitarsis		
17	Acanthomorpha ramulifera		
18	Spinidinium cf. spinidinium		
19	Spinidinium cf. spinidinium		
20	Polyosphaeridium subtile		
21	Chatangiella granulifera		
22	Alterbia acutula		
23	Isabelidium cooksoni		
24	Cribroperidium spp.		
25	Botryococcus spp.		
26	Trithrodinium suspactum		
27	Grassosphaera spp.		
28	Grassosphaera		
29	Hystriochloa staliatum		
30	Labridocystis chlamydata		
31	Odontochitina costata		
32	Chatangiella spp.		
33	Spiniferites "magnifica"		
34			
35			
36			
37			
38			
39	Bisaccates		
40	Callialasporites spp.		
41	Aquilapollenites spp.		
42	Cerebropollenites mesozoicus		
43	Loranthacites spp.		
44	Striatia bisaccata		
45	Baculatisporites/Osmundactites group		

WELL: - 30/3-3
CHART: - 2b of 5
FIGURE: - 4

CLIENT: - STATOIL
DEPTH: - 2200m-2500m
LOCATION: - NORWEGIAN NORTH SEA

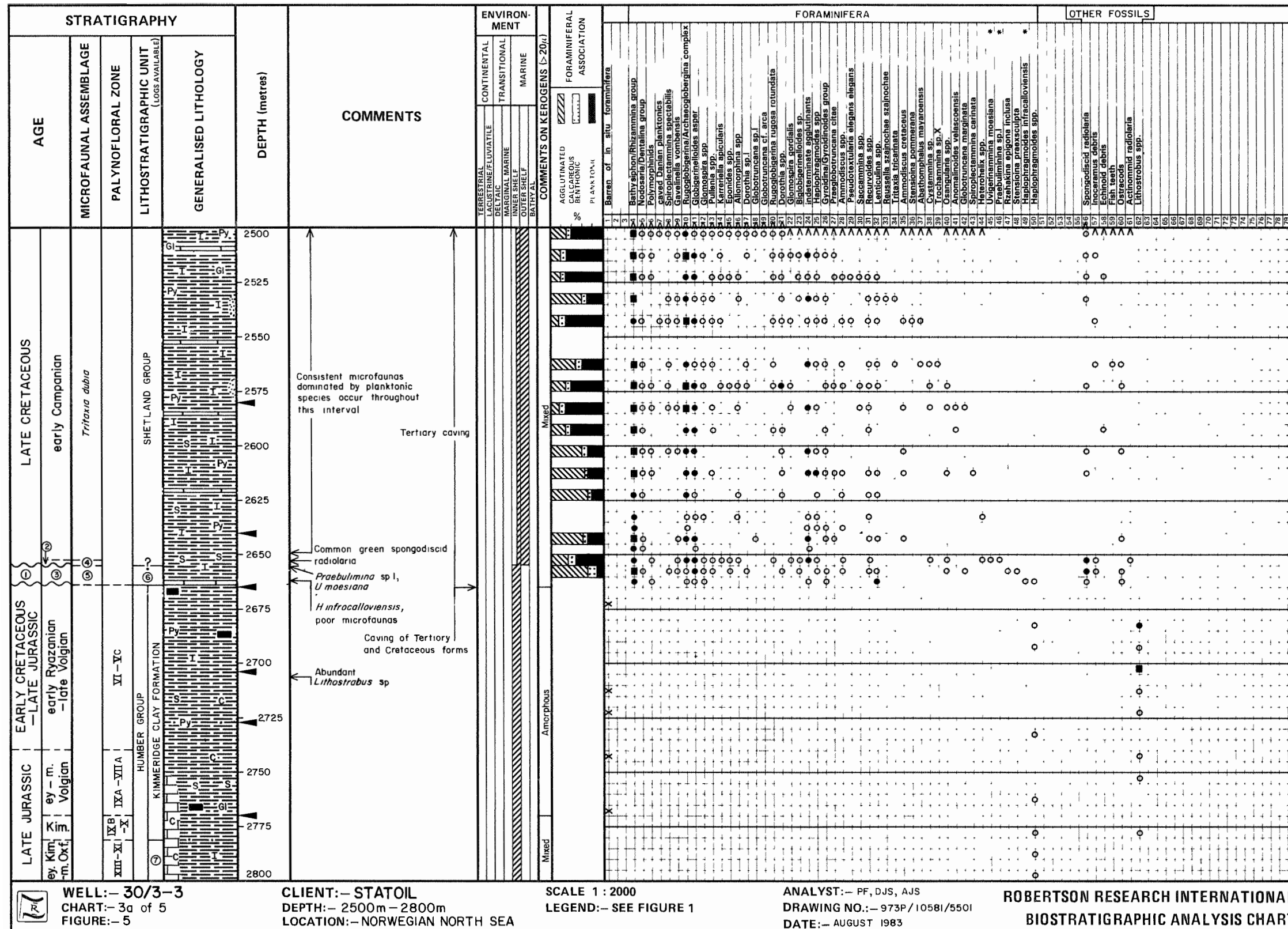
SCALE 1 : 2000
LEGEND: - SEE FIGURE 1

ANALYST: - JPGF, NHM
DRAWING NO.: - 973P/10581/5501
DATE: - AUGUST 1983

ROBERTSON RESEARCH INTERNATIONAL
BIOSTRATIGRAPHIC ANALYSIS CHART

Palynological analysis starts at 2500m

A acutula,
R fucatum,
S deltiense



WELL: - 30/3-3
 CHART: - 3a of 5
 FIGURE: - 5

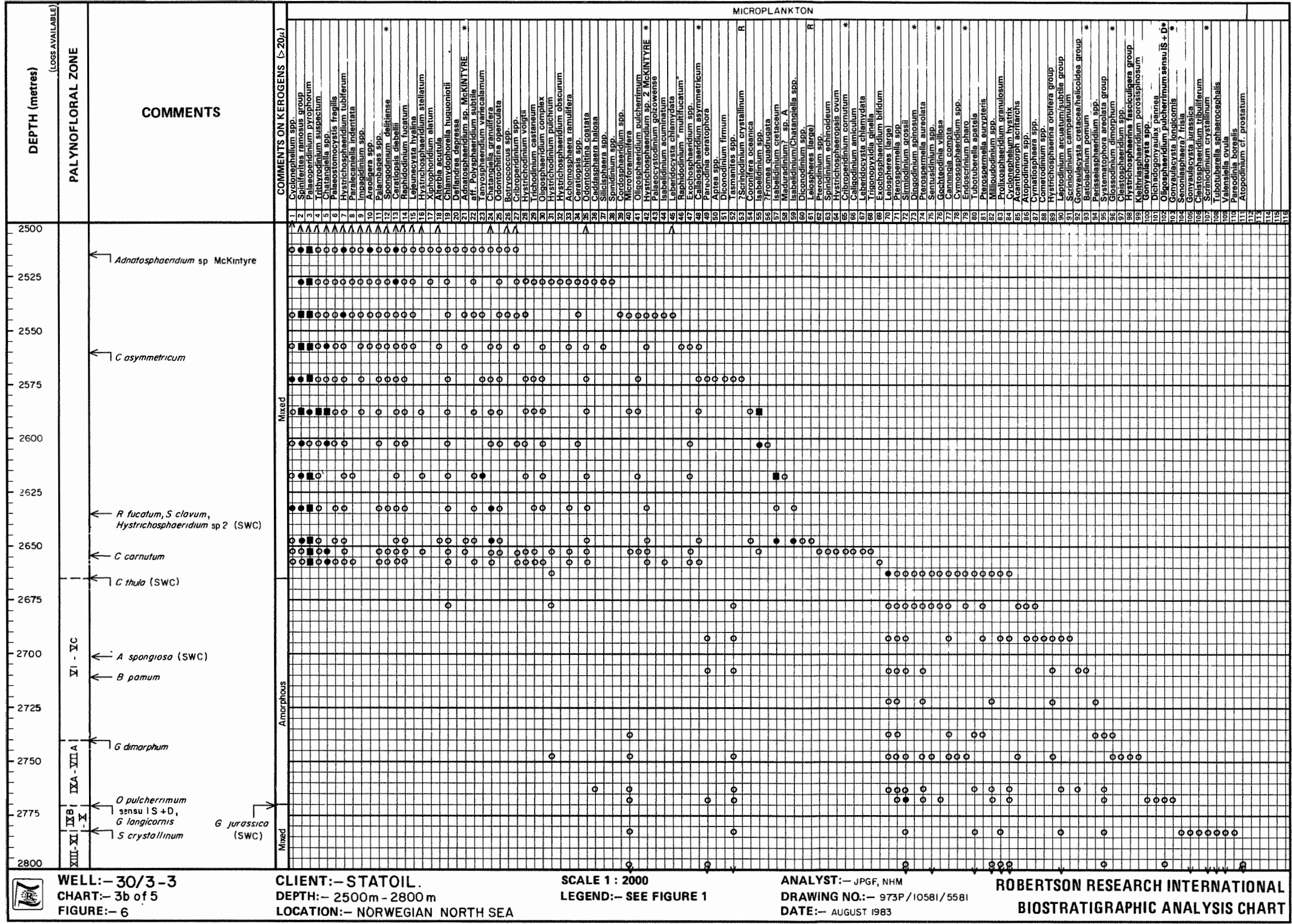
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 DEPTH: - 2500m - 2800m
 LOCATION: - NORWEGIAN NORTH SEA

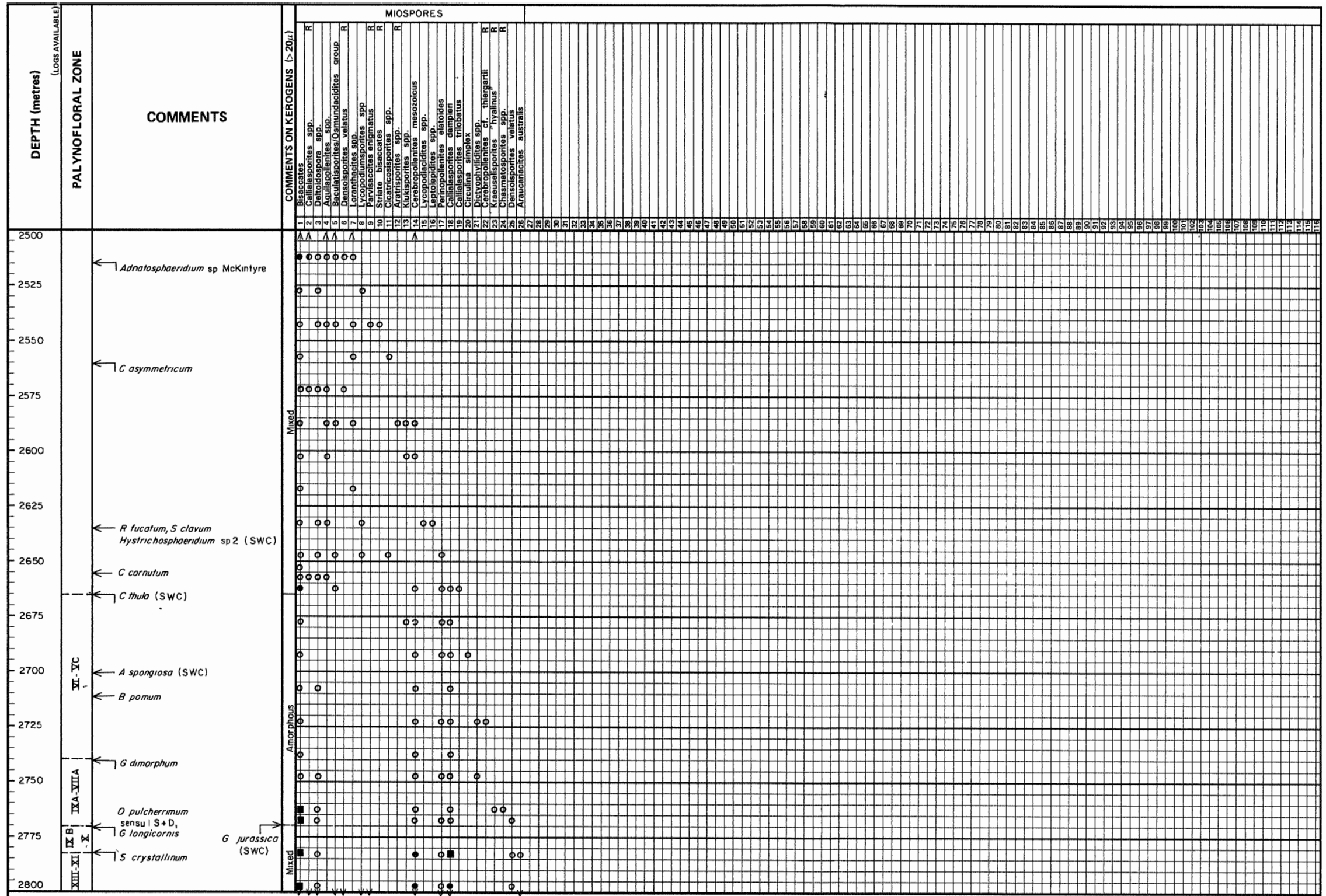
SCALE 1 : 2000
 LEGEND: - SEE FIGURE 1

ANALYST: - PF, DJS, AJS
 DRAWING NO.: - 973P/10581/5501
 DATE: - AUGUST 1983

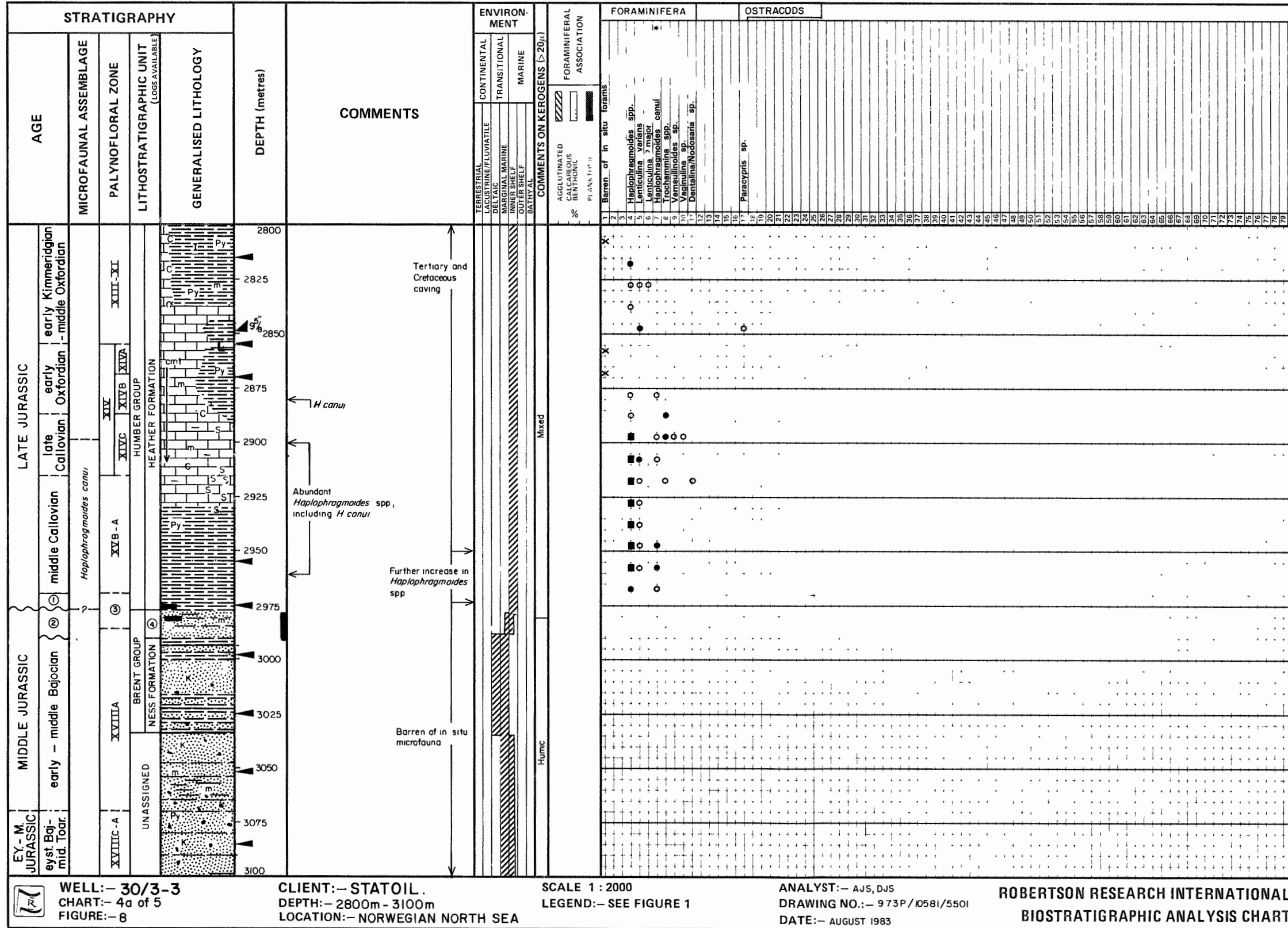
ROBERTSON RESEARCH INTERNATIONAL
 BIOSTRATIGRAPHIC ANALYSIS CHART

- ① EARLY CRETACEOUS
- ② Santonian
- ③ middle Albian - Barremian
- ④ *Stensioina*
- ⑤ *Recurvoides* / *Glomospira* - Calcareous benthonics / Ostracods
- ⑥ CROMER KNOLL GROUP
- ⑦ HEATHER FORMATION



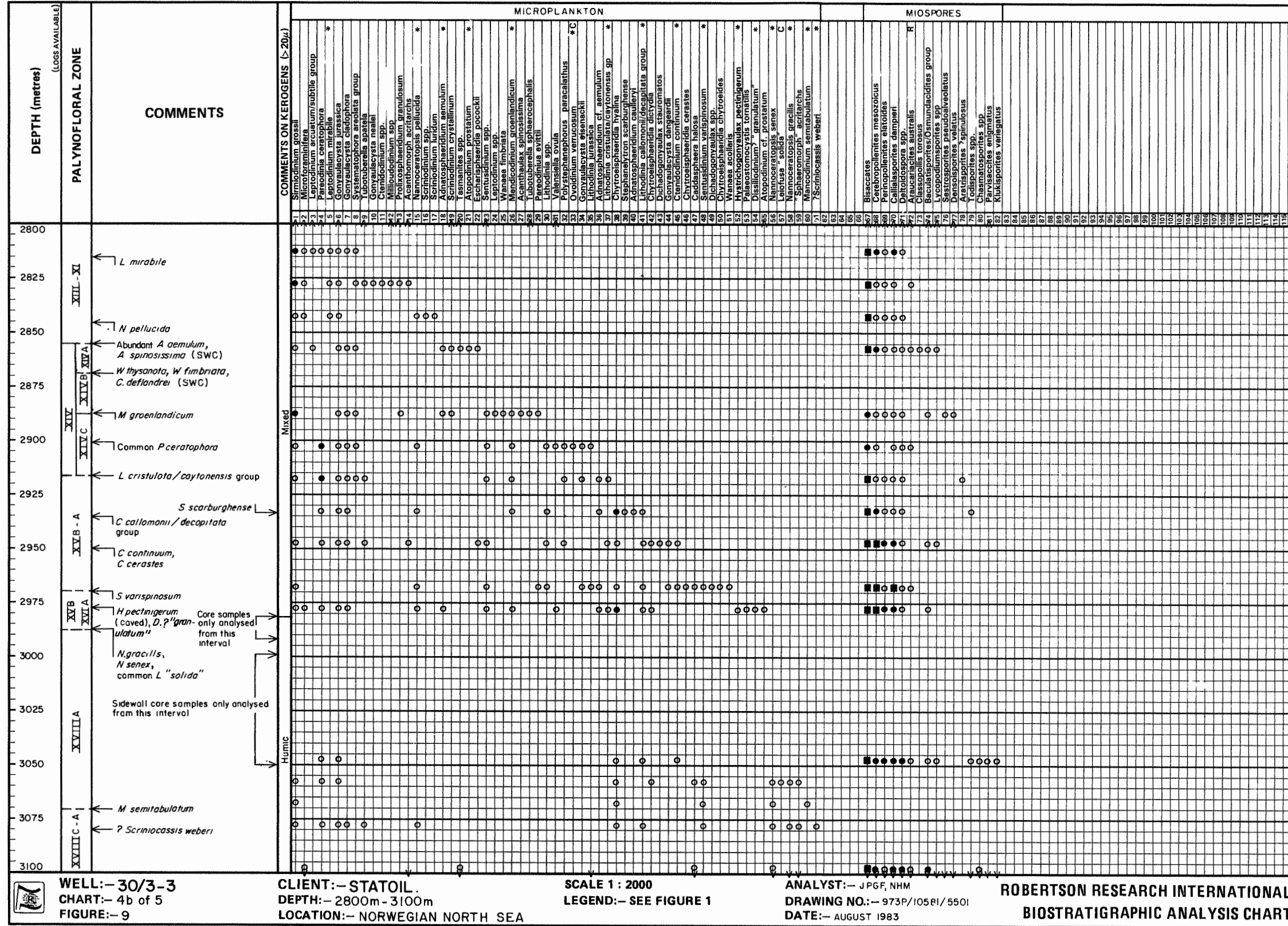


WELL: - 30/3-3
CLIENT: - STATOIL
CHART: - 3c of 5
DEPTH: - 2500m-2800m
FIGURE: - 7
LOCATION: - NORWEGIAN NORTH SEA
SCALE 1 : 2000
LEGEND: - SEE FIGURE 1
ANALYST: - JPGF, NHM
DRAWING NO.: - 973P/10581/5501
DATE: - AUGUST 1983
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BIOSTRATIGRAPHIC ANALYSIS CHART



- ① early Callovian
- ② early Bathonian - late Bajocian
- ③ XVIII - XVII
- ④ TARBERT FORMATION

WELL: - 30/3-3
CLIENT: - STATOIL.
CHART: - 4a of 5
DEPTH: - 2800m - 3100m
FIGURE: - 8
LOCATION: - NORWEGIAN NORTH SEA
SCALE 1 : 2000
LEGEND: - SEE FIGURE 1
ANALYST: - AJS, DJS
DRAWING NO.: - 9 73 P / 10561/5501
DATE: - AUGUST 1983
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BIOSTRATIGRAPHIC ANALYSIS CHART



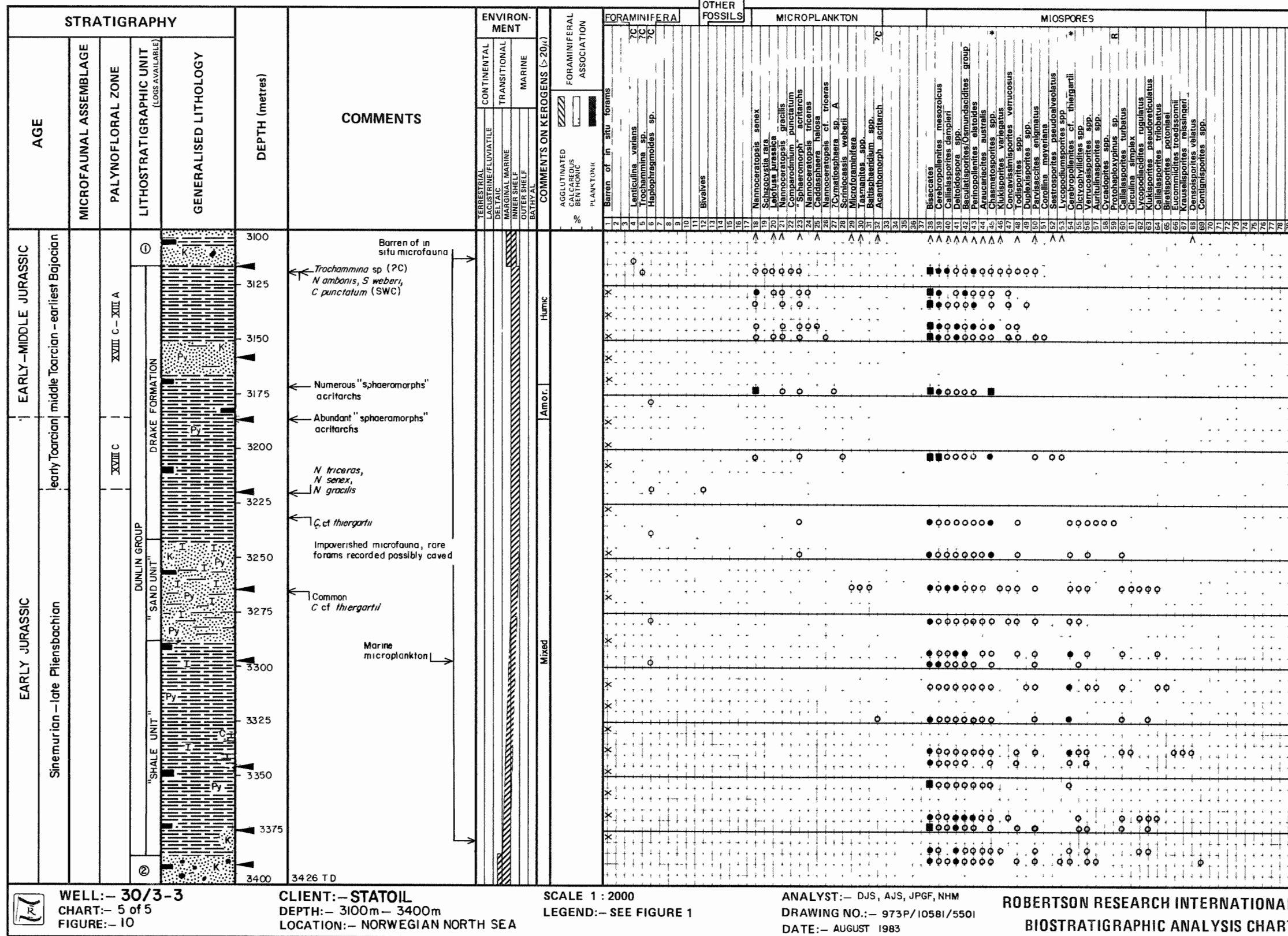
WELL:- 30/3-3
CHART:- 4b of 5
FIGURE:- 9

CLIENT:- STATOIL
DEPTH:- 2800m-3100m
LOCATION:- NORWEGIAN NORTH SEA

SCALE 1 : 2000
LEGEND:- SEE FIGURE 1

ANALYST:- J P G F, NHM
DRAWING NO.:- 973P/105R1/5501
DATE:- AUGUST 1983

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BIOSTRATIGRAPHIC ANALYSIS CHART



① UNASSIGNED
② STATFJORD FORMATION

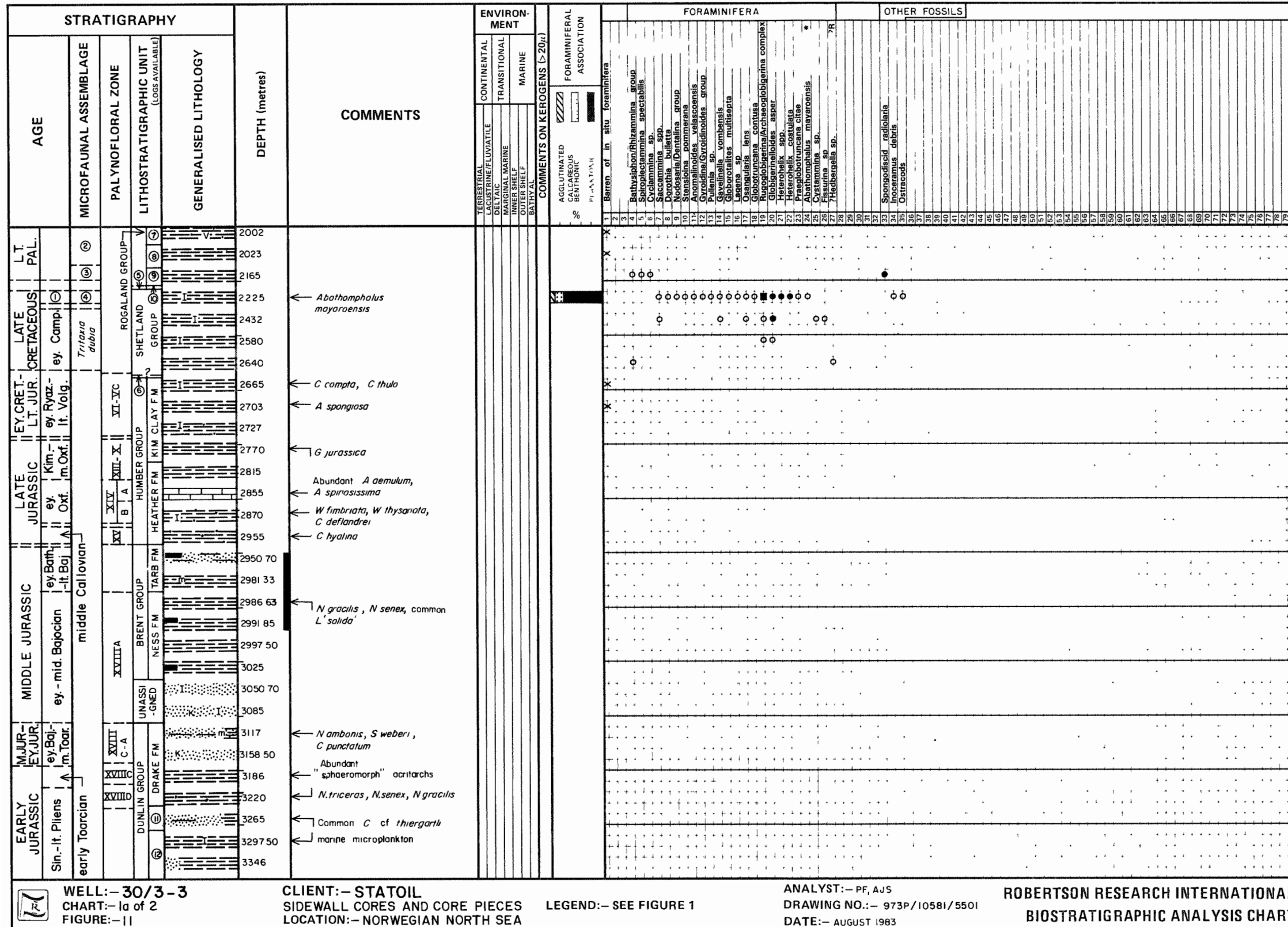
WELL: - 30/3-3
CHART: - 5 of 5
FIGURE: - 10

CLIENT: - STATOIL
DEPTH: - 3100m - 3400m
LOCATION: - NORWEGIAN NORTH SEA

SCALE 1 : 2000
LEGEND: - SEE FIGURE 1

ANALYST: - DJS, AJS, JPGF, NHM
DRAWING NO.: - 973P/10581/5501
DATE: - AUGUST 1983

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BIOSTRATIGRAPHIC ANALYSIS CHART



- ① late Maastrichtian - Maastrichtian
- ② *Cosciodiscus* sp I
- ③ *Trochammina* sp A
- ④ *Psudotextularia* - unnamed
- ⑤ MONTROSE GROUP
- ⑥ CROMER KNOLL GROUP
- ⑦ BALDER FORMATION
- ⑧ SELE FORMATION
- ⑨ LISTA FORMATION
- ⑩ MAUREEN FORMATION EQUIVALENT
- ⑪ " SAND UNIT "
- ⑫ " SHALE UNIT "

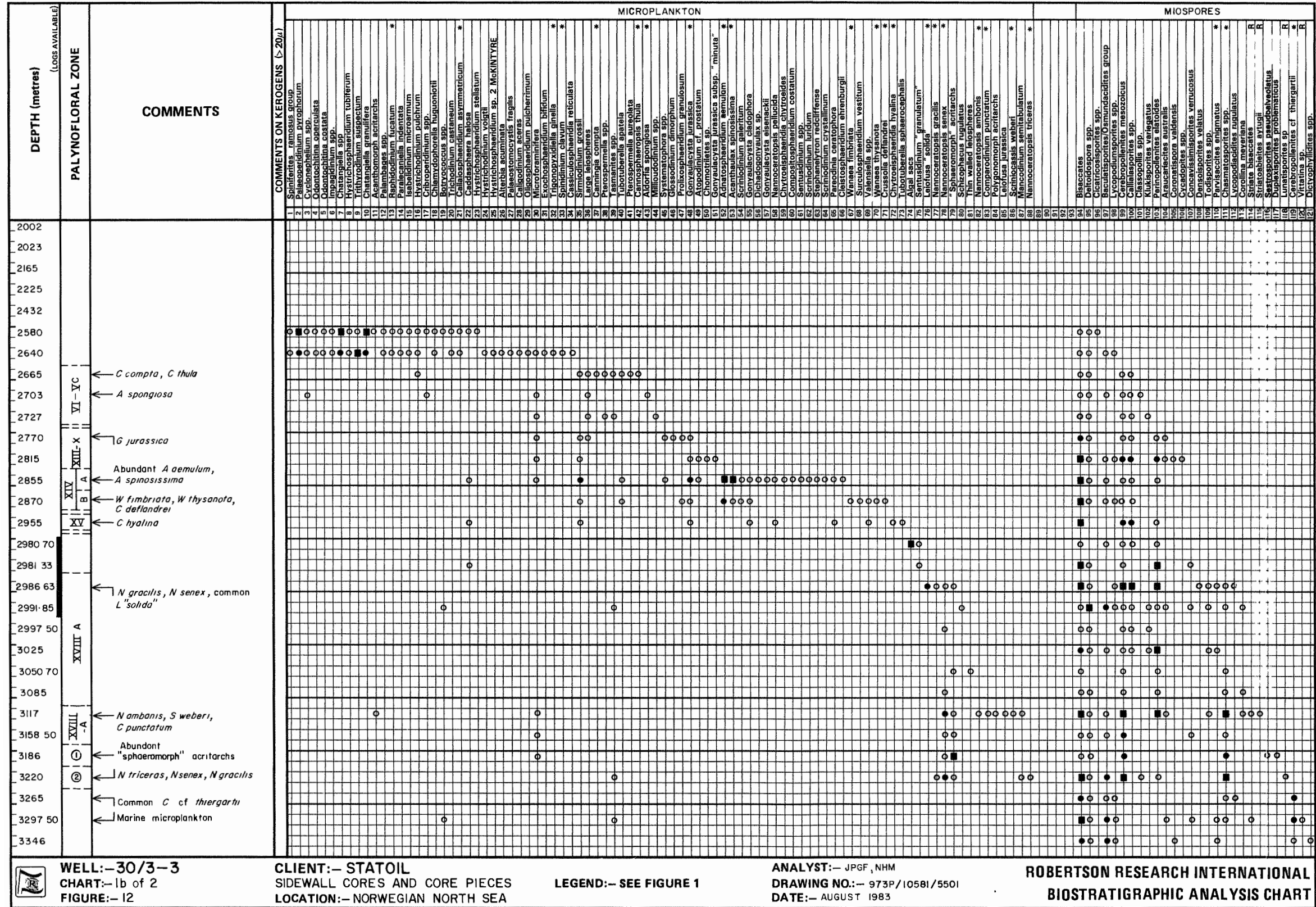
WELL: - 30/3-3
 CHART: - 1a of 2
 FIGURE: - 11

CLIENT: - STATOIL
 SIDEWALL CORES AND CORE PIECES
 LOCATION: - NORWEGIAN NORTH SEA

LEGEND: - SEE FIGURE 1

ANALYST: - PF, AJS
 DRAWING NO.: - 973P/10581/5501
 DATE: - AUGUST 1983

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 BIOSTRATIGRAPHIC ANALYSIS CHART



① XVIIIc
② XVIIIo

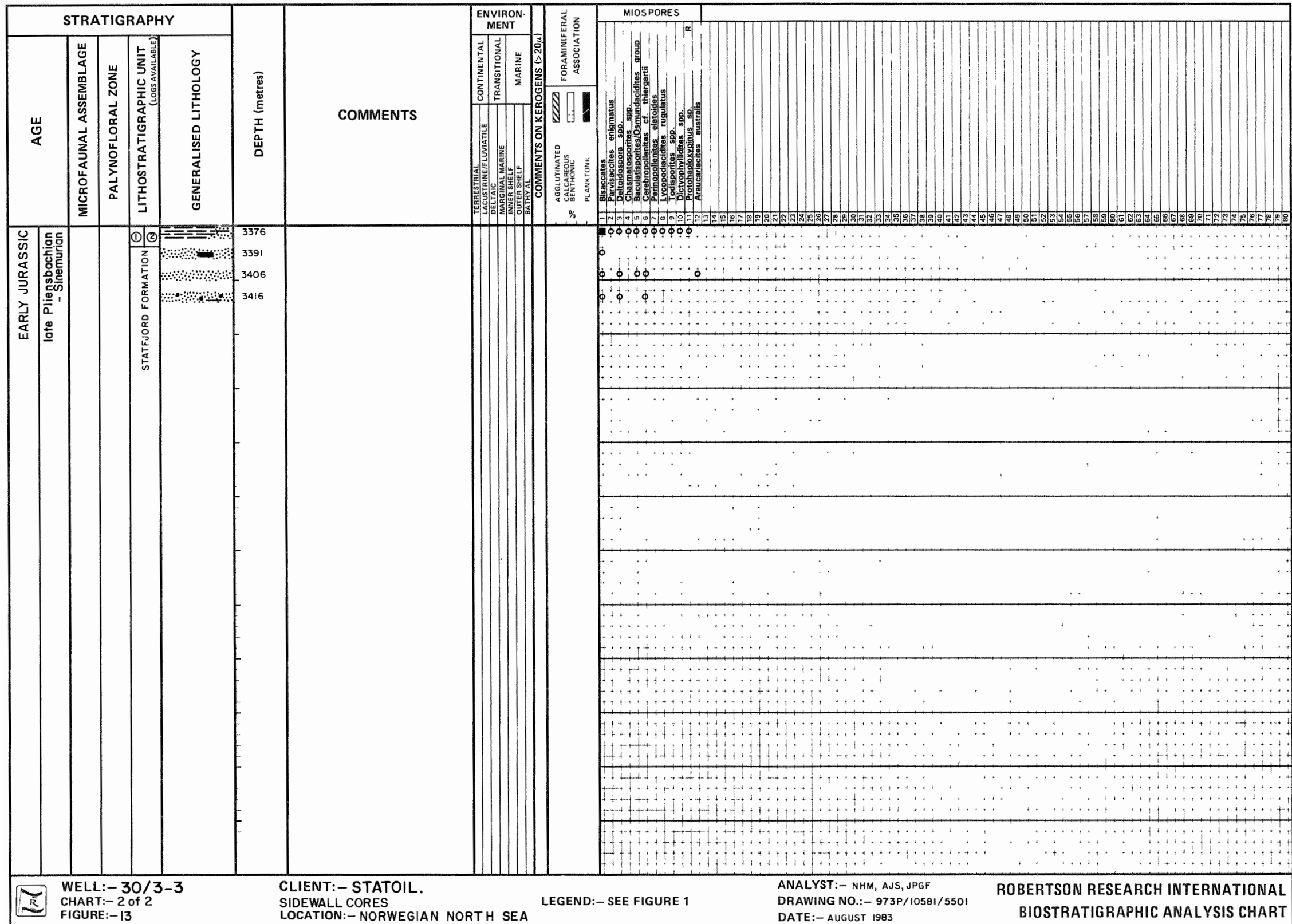
WELL: -30/3-3
CHART: -1b of 2
FIGURE: -12

CLIENT: - STATOIL
SIDEWALL CORES AND CORE PIECES
LOCATION: - NORWEGIAN NORTH SEA

LEGEND: - SEE FIGURE 1

ANALYST: - JPGF, NHM
DRAWING NO.: - 973P/10581/5501
DATE: - AUGUST 1983

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BIOSTRATIGRAPHIC ANALYSIS CHART



WELL:- 30/3-3
 CHART:- 2 of 2
 FIGURE:- 13

CLIENT:- STATOIL.
 SIDEWALL CORES
 LOCATION:- NORWEGIAN NORTH SEA

LEGEND:- SEE FIGURE 1

ANALYST:- NHM, AJS, JPGF
 DRAWING NO.:- 973P/10581/5501
 DATE:- AUGUST 1983

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① DUNLIN GROUP
 ② "SHALE UNIT"

STATOIL 30/3-3



ROBERTSON RESEARCH INTERNATIONAL

AREA NORWEGIAN NORTH SEA

SPUD DATE: 2nd APRIL 1983

COMPANY STATOIL

COMPLETION DATE: 31st MAY 1983

INTERVAL STUDIED: 1900m - 3418m

T.D. 3426m

Drawn by: 9739/0589/2502

ANALYST: J.P.F. FR. N.H.M. A.J.S. D.J.S.

Date: September 1983

