

APPENDIX III

BIOSTRATIGRAPHY OF WELL 31/2-2 (NORSKE SHELL)

by

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B. Prins

(SIPM, EP/12)

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1. INTRODUCTION

At request of Norske Shell, SIPM (EP/12.1) carried out a biostratigraphic study of the sedimentary sequence penetrated by well 31/2-2.

Previous experience in Mesozoic-Cenozoic stratigraphy of the general North Sea region suggested the following approach to this study :

- a) Tertiary : microfaunal analyses
- b) Cretaceous : calcareous nannoplankton investigations, supplemented by palynological work on the Lower Cretaceous, and
- c) Jurassic-Triassic : study of palynomorphs

Some overlap of investigations occurred across the boundaries of the major systems, e.g. microfaunal and nannoplankton studies on the basal Tertiary and higher Cretaceous strata.

The major objective of this well was the Jurassic sequence. For the Tertiary the main interest was on establishing the most important time-stratigraphic boundaries. Therefore, only ditch cuttings were available down to 1179 m, whilst SWS and cores in addition to cutting samples were recovered from 1185.5 m down to T.D.

The microfaunal and palynological data are recorded on separate distribution charts (encls. 1,2); the nannoplankton work is documented on text-figure 1.

The biostratigraphic results can be summarized as follows (not adjusted to wireline log readings) :

540	-	600	m	:	Miocene and/or younger
640	-	880	m	:	Oligocene
890	-	1217.5	m	:	Eocene
1224.9	-	1377	m	:	Paleocene
1379.6	m			:	Early-Middle Turonian
1381.0	-	1393.4	m	:	Early Turonian
1396.0	-	1421.9	m	:	Late Cenomanian
1427.6	-	1430.9	m	:	Late Albian
1434.0	-	1438.0	m	:	Hauterivian
1439.9	-	1446.5	m	:	Middle-Late Valanginian
1449.1	-	1454.9	m	:	(? Early) Valanginian
1458.5	-	1467	m	:	Early Valanginian
1470	m			:	Indeterminable
1472	m			:	Berriasian-Early Valanginian
1474	-	1489.5	m	:	Berriasian
1493.5	m			:	Portlandian-Berriasian
1495.5	-	1515.12	m	:	Portlandian
1521.07	-	1562.95	m	:	Kimmeridgian
1566.75	m			:	Indeterminable
1569	-	1600	m	:	Late Oxfordian
1604.24	-	1626.76	m	:	Early Oxfordian
1633.55	-	1649	m	:	Indeterminable
1657.5	-	1784	m	:	Middle Callovian
1793	m			:	Indeterminable
1812.5	-	1845.1	m	:	Late Bathonian-Early Callovian
1862	-	1982.5	m	:	Middle Bathonian
1998	-	2019	m	:	Indeterminable
2036	-	2140	m	:	Toarcian
2159	-	2181.5	m	:	Late Pliensbachian
2224.5	-	2340	m	:	(Early?) Pliensbachian
2351	-	2387	m	:	Hettangian-Sinemurian
2389	-	2571.5	m	:	Indeterminable

2. MICROFAUNA

2.1. General

The microfaunal studies concentrated on establishing the major age subdivision of the Tertiary interval. Consequently, the 14 sidewall cores and 47 cutting samples, selected for investigation, were primarily checked on marker types and no detailed studies were undertaken. Hence, the data plotted on the microfaunal distribution chart (encl. 1) do not represent the full picture of the faunal assemblages but only marker types and commonly occurring species.

The depths quoted in this chapter are sample depth figures, no adjustments to wireline log depths have been attempted. Especially in cases where only cutting samples were available, discrepancies may be unavoidable.

2.2. Miocene and/or younger (540-600 m)

The encountered assemblages consist essentially of species known from younger Tertiary-Quaternary strata, e.g.

Bulimina marginata D'ORBIGNY
Cibicides grossa TEN DAM & REINHOLD
Elphidiella hannai (CUSHMAN & GRANT)

Species restricted to the Miocene were not encountered. Therefore, only an age indication as Miocene and/or younger is possible. However, the occurrence of Elphidiella hannai in only sample 540 m could indicate that the Miocene/Pliocene boundary lies at about that level.

Depositional environment : The relatively poor benthonic fauna (consisting of calcareous types) and the scattered occurrences of single planktonics is considered indicative for an open-marine, inner neritic environment.

2.3. Oligocene (640-880 m)

This interval is characterized by poor to very poor assemblages, composed of (scarce) calcareous benthonic foraminifera, radiolarians, sponge spicules and echinoid remains.

The boundaries of this interval are defined by the highest occurrence of Turrilina alsatica ANDREAE (640 m) and the supposed top of Eocene strata (890 m) - see below.

Depositional environment : Based on the very scanty microfaunal information a general setting of open marine, shallow shelf is envisaged. Some samples were barren of foraminifera but showed relatively high amounts of lignite, thus possibly indicative for coastal influences.

2.4. Eocene (890-1217.5 m)

The top of this interval, at 890 m, is defined by a drastic change in the composition of the faunal assemblages, as well as in preservation (colour change).

The upper section (890-950 m) shows an occurrence of white-coloured, arenaceous types, such as Bathysiphon, Cyclamina placenta (REUSS) and Recurvoides obsoletum (GOES) in association with abundant radiolarians.

The middle part (960-1120 m) yielded relatively rich and diversified arenaceous assemblages. In addition to the already mentioned types also species as Cyclammina challinori HAYNES and C. amplectens GRZYBOWSKI are characteristic, together with the important markers Dorothia eocenica (CUSHMAN) and Textularia plummerae LALICKER, as well as Ammodiscus. The occurrence of pink sedimentary particles in cutting samples of 1100-1130 m together with Globigerina triloculinoides PLUMMER is considered significant as a similar phenomenon was observed in well 31/2-1 at the depth of 1170-1180 m.

The lowermost part of the Eocene interval (1130-1217.5 m) is characterized by the dominance of radiolarians and pyritized diatoms, such as

Coscinodiscus sp. 1 BETTENSTAEDT
C. sp. 2 BETTENSTAEDT
Triceratium sp. 1 BETTENSTAEDT

together with small numbers of arenaceous foraminifera.

Depositional environment : Outer neritic-bathyal.

2.5. Paleocene (1224.9 - 1377 m)

The Paleocene-Eocene boundary is taken to be located between SWS 1217.5 m (deepest occurrence of tuffaceous lithologies) and SWS 1224.9 m. The latter sample is devoid of microfaunal elements, a common feature for the topmost Paleocene strata in the general North Sea region.

The samples of the next interval (1233.5-1377 m) contain the following Paleocene marker species :

Rzehakina minima CUSHMAN & RENZ
Saccammina rhumbleri (FRANKE)
Trochammina ruthven-murrayi CUSHMAN & RENZ

The major part of the foraminiferal fauna, however, consists of arenaceous types also occurring in the overlying Eocene interval, e.g.

Spiroplectammina spectabilis (GRZYBOWSKI)
Glomospira charoides (JONES & PARKER)
Haplophragmoides kirki WICKENDEN
Cyclammina chalinori HAYNES

Conspicuous is the green colour of parts of the faunas in the higher part of this section, a phenomenon regionally known to be characteristic for younger Paleocene deposits.

From 1284.3 m downwards very coarse-grained Bathysiphon specimens were encountered.

Cutting sample 1371 m yielded, amongst others, Ceratobulimina tuberculata BROTZEN, indicative for Lower Paleocene.

SWS 1372 m and cutting sample 1377 m yielded Lower Tertiary faunas, but on log evidence the Paleocene-Cretaceous boundary is located at 1374 m.

Depositional environment : Outer neritic-bathyal.

2.6. Cretaceous (1383-1398 m (deepest sample examined))

The sidewall samples taken at 1379.6, 1381.7 and 1384.4 m contained only a few undiagnostic arenaceous foraminifera. On the other hand, the cutting samples 1383, 1389, 1395 and 1398 m yielded, apart from rich Paleocene faunas, also some Upper Cretaceous types, as

Guembelina globulosa (EHRENBERG)
Hedbergella debrioensis (CARSEY)
H. planispira (TAPPAN)
H. ? hoelzli (HAGN & ZEIL)

indicative for a Cenomanian-Turonian age.

SWS 1388.9 m yielded a mainly benthonic fauna with as the most important components :

Planularia bradyana (CHAPMAN)
Textularia bettenstaedti BARTENSTEIN & OERTLI
Verneuilinoides subfiliformis BARTENSTEIN
Trochammina concava CHAPMAN

and only a few specimens of Praeglobotruncana.

Based on the above fauna, an Albian-Cenomanian age is concluded. This evidence is contradictory to the available nannoplankton data, which conclude a Turonian age. The last dating is preferred as most ranges of Cretaceous benthonic foraminifera are, as yet, not well established for the northern North Sea region.

SWS 1396 and 1398 m are characterized by a fauna entirely composed of planktonic foraminifera, e.g.

Globigerinoides alvarezii (ETERNOD OLVERA)
Hedbergella simplex (MORROW)
H. debrioensis (CARSEY)

indicative for a Late Albian-Coniacian age.

Depositional environment : Open marine, outer neritic-bathyal.

3. CALCAREOUS NANNOPLANKTON

3.1. General

The investigation of the nannofloras in well 31/2-2 was confined to the interval 1180.5-1501.5 m, mainly representing the Cretaceous interval. A total of 53 sidewall cores have been prepared and analysed.

The species composition of the Cretaceous nannofloras differs considerably from other localities reported in literature and/or known from own investigations. However, a detailed biostratigraphic subdivision could be obtained by the application of evolutionary trends in a few groups of nannoplankton, in combination with rare occurrences of important marker species. This enabled the recognition of several gaps of differing magnitude, indicating the following Cretaceous intervals as absent : Barremian-Lower/Middle Albian, Early-Middle Cenomanian, Late Turonian-Maastrichtian.

As in the previous chapter, quoted depth figures are sample depths. No adjustments to wireline log depths have been introduced.

3.2. Early Tertiary (1180.5-1372.0 m)

All investigated samples were devoid of nannofossils. This interval may represent the (? decalcified) Middle/Late Paleocene-Eocene shales, widely distributed in the North Sea Basin.

No indications have been found for the presence of Early Paleocene (Danian) deposits.

3.3. Turonian (1379.6-1393.4 m)

1379.6-1393.4 m : Early to Middle Turonian, zone NK 11.

The presence of seven and eight rayed specimens in the Eprolithus lineage, the fairly large size and frequent occurrence of Prediscosphaera intercisa, the decreasing number of P. ponticula, the absence of P. columnata and Podorhabdus albianus indicate this interval as not older than Turonian. Furthermore, the presence in all samples of Zeugrhabdotus theta (a species becoming extinct halfway the Turonian) restrict the age of this interval to Early to Middle Turonian, zone NK 11.

3.4. Cenomanian (1396,0-1421,9 m)

1396,0-1421,9 m : Late Cenomanian, zone NK 10.

The age-diagnostic criteria are the co-occurrences of Podorhabdus albianus, Lithraphidites acutum, Prediscosphaera columnata and large specimens of P. intercisa.

The rather sudden change in the floral composition of the Eprolithus and Prediscosphaera lineages between the lowest Turonian sample at 1393,4 and the highest Cenomanian sample at 1396,0 m points to a small hiatus, which may coincide with a distinct change in log expression at around 1395 m.

3.5. Albian (1427,6-1430,9 m)

1427,6-1428,8 m : Late Albian, subzone NK 9a.

This age determination is based on the co-occurrence of Nannofossil sp. 138, Rucinolithus irregularis and Eiffelolithus turriseiffeli.

A distinct hiatus exists between the Late Cenomanian sample at 1421,9 m and the highest Albian sample at 1427,6 m, the two subzones NK 9b and 9c being absent. The extent of the hiatus is demonstrated by a marked difference in the Prediscosphaera assemblages above and below the hiatus, which might be located at 1424 m where a distinct change in log expression occurs.

1430,9 m : Albian ?, zones NK 8 - NK 9a ?.

The very poor preservation of the nannoflora in this sample does not allow a more precise age determination. The presence of a few specimens of Prediscosphaera columnata may indicate this sample not being older than Early Albian, subzone NK 8a. However, their presence may also be explained by mud penetration into the sidewall core, a phenomenon regularly observed in the samples from this well.

Note : The three Albian sidewall cores of this well consisted lithologically of various parts with markedly different colours. From the sample at 1427,6 m several chips with different colours have been prepared separately, but the nannofloras of all chips were found to be identical.

3.6. Hauterivian (1434.0-1438.0 m)

1434.0-1438.0 m : Huaterivian, zone NK 4.

The presence of Tegumentum striatum suggests a Hauterivian age for this interval. The absence of Palaeopontosphaera salebrosa in the upper two sidewall cores at 1434.0 and 1435.5 m and its common occurrence at 1438.0 m may indicate that the upper two samples belong to the upper part of zone NK 4 and the lower sample to the lower part of this zone.

3.7. Valanginian (1439.9-1454.9 m)

1439.9-1446.5 m : Middle to Late Valanginian, zone NK 3.

The age of these samples has been based on the presence of the very distinct species Micrantholithus speetonensis.

1449.1-1454.9 m : (Early ?) Valanginian, zone NK 2 upper - NK 3.

The presence of several specimens of Nannoconus bermudezii in both samples indicates this interval not being older than Valanginian. The complete absence of Micrantholithus or allied genera in these (and two deeper samples) suggests that the absence of M. speetonensis may also be explained by unfavourable environmental conditions. Therefore, a Middle to Late Valanginian age for these two samples is not excluded.

3.8. Berriasian-Valanginian (1457.5-1474.0 m)

1457.5 m : Late Berriasian to Valanginian, zones NK 2-3.

Although this sample contains a rich and diversified nannoflora, no zonal indicators could be found, even after detailed analysis of the flora. Therefore, only an age-range can be given based on the ages of over- and underlying samples.

1458.5-1470.0 m : Late Berriasian to Early Valanginian, zone NK 2

The age of this interval has been based on the presence of Brachyolithus quadratus and the allied Nannofossil sp. B.

1472.0-1474.0 m : Early Berriasian to Early Valanginian, zones NK1-2.

The nannoflora is badly preserved and poor in species, the prevailing type being Watznaueria barnesae. Its presence, combined with the absence of Ellipsagelosphaera suggests these samples to be not older than Berriasian.

3.9. Berriasian and/or older (1482,0-1493,5 m)

The majority of the ten samples, covering the section of 1479.0-1501.5 m, did not contain any nannofossils. Only 3 samples, between 1482.0 and 1493.5 m, yielded a few to rare specimens of Ellipsagelosphaera indicative for an age range of Bajocian-Early Berriasian.

4. PALYNOLOGY

4.1. General

A total of 99 samples have been examined, viz. 19 core samples, 78 sidewall cores and 2 cutting samples. All samples have been prepared following the standard preparation method, consisting of treatment with HCl and HF, followed by a heavy liquid separation (zincbromide with S.G. of 2.2). The organic residue was finally sieved through microsieves of 15 microns, to concentrate palynomorphs in general, and of 30 microns to concentrate dinocysts. Especially in the Jurassic interval, the organic residues received an oxidation by means of cold HNO₃, prior to sieving, in order to gain a further concentration of dinocysts.

For the major part of the Jurassic sequence only dinocysts have been recorded. General evidence has shown this group to be far more suitable for a subdivision of the Middle-Upper Jurassic than sporomorphs, provided marine strata are investigated. However, the often sandy development of these beds influenced dinocysts recoveries unfavourably. In the Pliensbachian-uppermost Triassic interval, dinocyst diversities decrease drastically, in general agreement with the known evolutionary development of the group. Therefore, selected sporomorphs were applied for a further subdivision of this part of the drilled sequence, starting from the top occurrence of the dinocyst species Nannoceratopsis gracilis.

The relative frequencies of the determined palynomorphs have been recorded on a distribution chart (encl. 2). A number of samples appear to yield fair amounts of palynomorphs obviously derived from caved material or mud infiltration.

Throughout the studied section the sporomorph colour was estimated as "upper light", corresponding to a FCC of less than approximately 69. The quoted depths are sample depths, no adjustments to wireline log depths have been attempted.

4.2. Cenomanian-Early Turonian (1381-1421.9 m)

Four sidewall cores were examined. These yielded rich and well-preserved dinocyst assemblages with a number of significant species, e.g.

Palaeohystrichophora infusorioides (Late Albian to Campanian)
Odontochitina costata (id.)
Phoberocysta ceratioides (id.)
Xiphophoridium alatum (Late Albian to Early Turonian)
Cyclonephelium hughesii (id.)
Cometodinium obscurum (Albian to Early Turonian)
Microdinium irregulare (Cenomanian to Maastrichtian)
Exochosphaeridium striolatum (Cenomanian and younger)

4.3. Late Albian-Cenomanian (1430.9 m)

One sidewall core yielded a few dinocysts only, with

Palaeohystrichophora infusorioides (Late Albian to Campanian).

4.4. Valanginian-Cenomanian (1435.5-1454.9 m)

From this interval six sidewall cores and one cutting sample were examined. Five samples proved to be barren of palynomorphs. The cutting sample, at 1440 m, yielded only Tertiary palynomorphs derived from mud contamination. One SWS, at 1446.5 m, contained, in addition to several obviously allochthonous palynomorphs, a few specimens of dinocysts which could well be autochthonous, viz.

Pareodinia ceratophora (top occurrence in Barremian)
Exochosphaeridium striolatum (base in Cenomanian)

As the co-occurrence of these species is conflicting, no conclusion on the age of this sample can be made.

4.5. Valanginian (1461.5-1467 m)

Two sidewall cores yielded a very rich, well-preserved Valanginian dinocyst flora with as typical species :

Heslertonia heslertonense (base in Valanginian)
Pseudoceratium pelliferum subsp. solocispinum (id.)
Muderongia simplex, circumvate variety (id.)
Achomosphaera neptuni (id.)
Occisucysta tentoria (id.)
Trichodinium ciliatum (acme in Valanginian)
Speetonia delicatula (Valanginian)
Endoscrinium pharo (top in Valanginian)
Occisucysta evittii (Berriasian to Valanginian)

4.6. Indeterminable (1470 m)

A cutting sample from 1470 m yielded only Tertiary palynomorphs, obviously derived from mud contamination, and the Berriasian to Valanginian age is based on the ages of over- and underlying samples.

4.7. Berriasian (1474-1489.5 m)

Three sidewall cores yielded rich and well-preserved dinocyst assemblages of Berriasian age, based on co-occurrence of :

Occisucysta evittii (Berriasian to Valanginian)
Gonyaulacysta sp. A DAVEY (top in Berriasian)
"Escharisphaeridia ryazanica" manuscript name (= Canningia sp.1)
(typical acme in Berriasian)
Parvocavatus spinosus (Portlandian to Berriasian)
Tanyosphaeridium sp. 1 (top in Berriasian)
Tubotuberella apatela (top regular occurrence in Berriasian)
Hystrichogonyaulax sp. 2 (id.)
"Cordosphaeridium daveyi" manuscript name (= Kleithria -
sphaeridium sp. A DAVEY) (top in Berriasian)
Systematophora cf. complicata (id.)
Leptodinium eumorphum (id.)

4.8. Portlandian (1495.5-1515.12 m)

From this interval, three SWS and one core sample have been studied. These yielded rich dinocyst assemblages with :

Millioudodinium mamilliferum (base occurrence in Portlandian)
Pareodinia dasyforma (id.)
Parvocavatus spinosus (id.)
Hystrichodinium pulchrum (base regular occurrence in
Portlandian)
Glossodinium dimorphum (Kimmeridgian to Portlandian)
Hystrichogonyaulax sp. 1 (typical acme in Portlandian)

4.9. Kimmeridgian (1521.07-1562.95 m)

Two SWS and 8 cores were examined. Below 1535 m a low diversity and qualitatively poor dinocyst flora is common.

A Kimmeridgian age was concluded on :

Hystrichogonyaulax cladophora (top(regular) occurrence in
Kimmeridgian)
Muderongia simplex - cornucavate variety (id.)
Sentusidinium rioultii (id.)
Leptodinium mirabile (id.)
Cassiculosphaeridia sp. 1 (base (regular) occurrence in
Kimmeridgian)
Systematophora cf. complicata (id.)
Glossodinium dimorphum (id.)
Senoniasphaera jurassica (id.)
Cyclonephelium distinctum (id.)
Hystrichodinium pulchrum (id.)
Acanthaulax sp. A IOANNIDES et al. (id.)

A core sample, at 1521.07 m, contains also a high number of

Gonyaulacysta longicornis s.s. (common and typical in
Kimmeridgian)

4.10. Indeterminable (1566.75 m)

One core sample, at 1566.75 m, proved to be barren of palynomorphs, and the Oxfordian-Kimmeridgian age is inferred from the ages of the over- and underlying samples.

4.11. Oxfordian (1569-1626,76 m)

A total of 8 core samples and 1 cutting sample have been examined. One core sample, at 1574.80 m, was barren of autochthonous dinocysts but contained numerous Tertiary and Late Cretaceous palynomorphs derived from mud infiltration into the core.

An Oxfordian age was concluded on :

Nov. gen. G sp. 1 (top and base occurrence in Oxfordian, very
rare in Callovian)
Endoscrinium galeritum (top in Oxfordian)
Adnatosphaeridium aemulum (id.)
Stephanelytron (id.)
Gonyaulacysta jurassica subsp. longicornis (id.)
G. areolata (id.)
Leptodinium mirabile (base in Oxfordian)
Scriniodinium crystallinum (acme in Oxfordian)
Prolixosphaeridium capitatum (id.)

The interval can be subdivided into :

Late Oxfordian (1569 -1600 m)
Early Oxfordian (1604.24-1626.76 m)

based on the absence and presence of Gonyaulacysta areolata, respectively.

4.12. Indeterminable (1633.55-1649 m)

One core sample, at 1633.55 m, proved to be barren of dinocysts. A sidewall core, at 1649 m, yielded a dinocyst assemblage which was too poor for age determination.

The Callovian-Early Oxfordian age could only be inferred from the stratigraphic position of the interval.

4.13. Middle Callovian (1657.5-1784 m)

Fourteen sidewall cores have been examined. These yielded rather poor to moderately rich dinocyst assemblages, in which the presence of the following species is significant :

Hystrichogonyaulax pectinigera (Middle Bathonian to Middle Callovian)
Gonyaulacysta jurassica (Latest Bathonian to Kimmeridgian)
Lithodinia jurassica (Middle to Late Callovian)
"Lithodinia suturocomplexa" manuscript name (= L. sp. 1) (Early to Middle Callovian)
Ctenidodinium ornatum (Callovian to Oxfordian)
Sentusidinium rioultii (base regular occurrence in Callovian)
Hystrichogonyaulax sp. 5 (Middle Bathonian to Middle Callovian)

4.14. Indeterminable (1793 m)

This sidewall sample contains a poor dinocyst flora, and the Callovian age is inferred.

4.15. Late Bathonian-Early Callovian (1812.5-1845.1 m)

Seven sidewall cores yielded moderately rich dinocyst assemblages. The age determination is based on the presence of

Hystrichogonyaulax pectinigera (id.) (Middle Bathonian-Middle Callovian)
Emmetrocyta sp. nov. (Late Bathonian to Early Callovian)
cf. Lunatadinium sp. 1 (Bathonian to Callovian)
Lithodinia sp. 6 (Bathonian to Callovian)
Lithodinia sp. 10 (so far only found in Late Bathonian to Early Callovian strata)

4.16. Middle Bathonian (1862-1982.5 m)

Seven sidewall samples have been analysed. The age is based on significant species such as

Hystrichogonyaulax sp. 5 (Middle Bathonian to Early Callovian)
Lunatadinium sp. 1 (Middle Bathonian to Middle Callovian)
Ctenidodinium sp. 1 (so far found only in Middle Bathonian)

and scattered occurrences of

Hystrichogonyaulax sp. 6 (so far found only in Middle Bathonian)
Pareodinia sp. 2 (id.)
Pareodinia sp. 4 (id.)

The sidewall core at 1982.5 m yielded very poor dinocyst assemblage. The presence of

Hystrichogonyaulax cladophora (base in Middle Bathonian)

is indicative for an age not older than Middle Bathonian, and the sample is included in this interval.

4.17. Indeterminable (1998-2019 m)

The sidewall cores at 1998 m and 2019 m proved to be barren of dinocysts, and the Toarcian to Middle Bathonian age has been inferred from the over- and underlying samples.

4.18. Toarcian (2036-2140 m)

Five sidewall cores were examined. These yielded relatively rich palynomorph assemblages, dominated by saccate sporomorphs and the dinocyst species Nannoceratopsis gracilis ("teardrop" variety). Significant sporomorph types are :

Klukisporites variegatus (base in Toarcian)
Chasmatosporites magnolioides (acme in Toarcian)
Cerebropollenites macroverrucosus (base in Sinemurian)

The consistent absence of the otherwise omnipresent sporomorph group of Callialasporites may suggest that only Early Toarcian is represented.

4.19. Late Pliensbachian (2159.0-2181.5 m)

Three sidewall cores from this interval yielded rich sporomorph assemblages. The presence in sidewall core 2181.5 m of

Nannoceratopsis gracilis (base in Late Pliensbachian)

suggests that the age of this interval is not older than Late Pliensbachian.

4.20. (Early ?)Pliensbachian (2224.5-2340.0 m)

Six sidewall cores from this interval have been examined. They yielded rich sporomorph assemblages, in which

Chasmatosporites thiergarti

shows a distinct acme throughout the interval, which is diagnostic for a Pliensbachian age.

Based on the consistent absence of

Nannoceratopsis gracilis (base in Late Pliensbachian)

the age of this interval may be limited to Early Pliensbachian.

4.21. Hettangian-Sinemurian (2351.0-2387.0 m)

Two sidewall cores in this interval yielded only a rather poor sporomorph assemblage, with rare occurrence of

Cerebropollenites mesozoicus (base regular occurrence in Sinemurian)

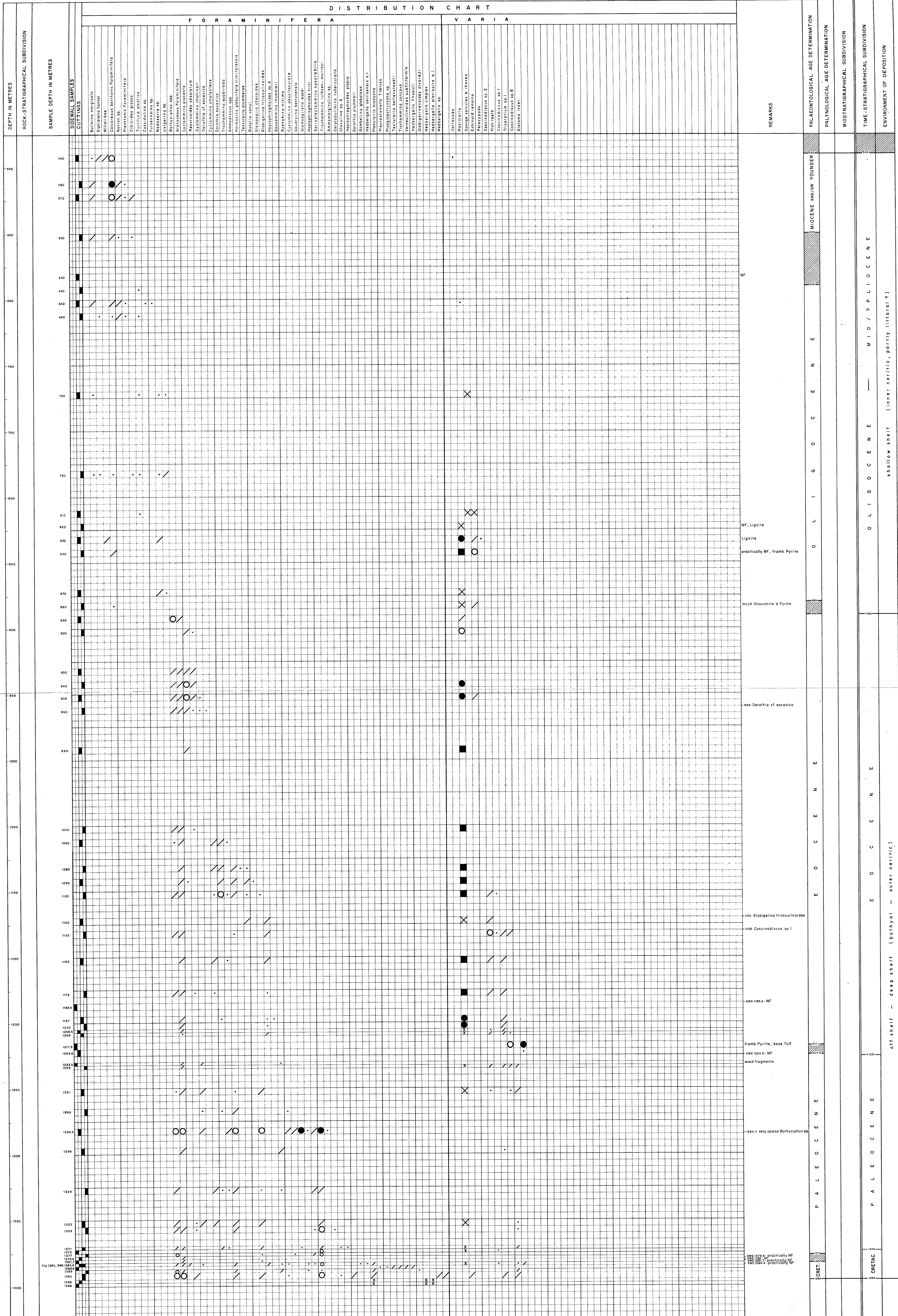
Quadraeculina anellaeformis (Rhaetian to Lower Jurassic)

Chasmatosporites magnolioides (Early to Middle Jurassic)

Based mainly on the absence of pre-Hettangian markers, as well as of absence of Pliensbachian elements, a Hettangian to Sinemurian age has been inferred to these two samples.

4.22. Indeterminable (2389.0-2571.5 m)

Eleven sidewall cores in this interval have been examined. They were all devoid of palynomorphs.



DEPTH IN METRES

ROCK-STRATIGRAPHICAL SUBDIVISION

SAMPLE DEPTH IN METRES

SIDEWALL SAMPLES CUTTINGS

DISTRIBUTION CHART

FORAMINIFERA

VARIATA

REMARKS

PALAEONTOLOGICAL AGE DETERMINATION

PALYNOLOGICAL AGE DETERMINATION

BIOSTRATIGRAPHICAL SUBDIVISION

TIME-STRATIGRAPHICAL SUBDIVISION

ENVIRONMENT OF DEPOSITION

LEGEND

- single occurrence
- 2-5 specimens
- 21-100 "
- >100 "
- X quantity not specified

NOTE: only marker types and common species are plotted.

SHELL INTERNATIONALE PETROLEUM MAATSCHAPPIJ NV
 THE HAGUE EXPLORATION & PRODUCTION

NORWAY OFFSHORE

MICROFAUNAL DISTRIBUTION CHART

NORSKE SHELL WELL 3/2-2

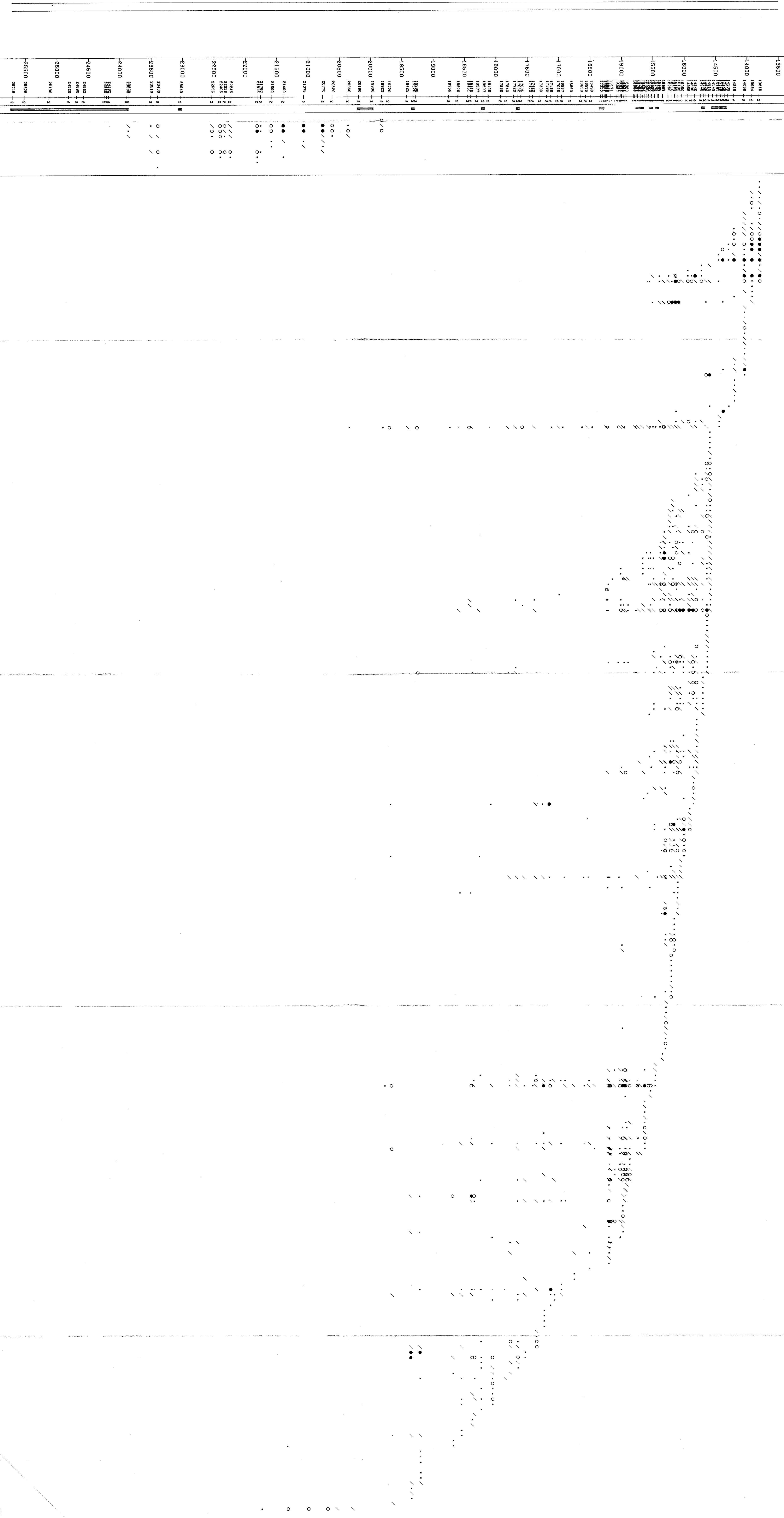
INTERVAL 540 m - 1398 m
 Scale 1:1000

Author: E. Noordermeer Date: July 1980
 Appendix II Well Return 3/2-2 Exec: 1 Draw No: 6.66227/1

DEPTH IN DECIMETRES

Table with columns: SAMPLES EXAMINED, GENUS, SPECIFIC DETERMINATION, INDEX. Lists various microfossil species and their corresponding index numbers.

REMARKS, BIOFACIES DEVELOPMENT, MICROFAUNAL ZONATION, MICROFLORAL ZONATION, TIME-STRATIGRAPHIC SUBDIVISION, DEPOSITIONAL ENVIRONMENTS.



STANDARD LEGEND, SAMPLES EXAMINED, SHELL INTERMEDIATE PETROLEUM N.V., NORWAY OFFSHORE, PALEONTOLOGICAL DISTRIBUTION CHART, NONSBE SHELL WELL 31/2-2, STRIKE 1:2500, SCALE, SHELL NO. 03124614, DATE: JUN 1980, SHELL NO. 03622712.