

SNEAP (P)

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LABORATOIRE DE GEOLOGIE DE BOUSSENS

-  
GEO/LAB Bss n° 7/1536 RP  
/fr

16/3-2 WELL (NORWAY)

SOURCE-ROCK IDENTIFICATION  
AND  
ORGANIC CATAGENESIS

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Reference : Order n° 033 128

R. CUSSEY : 7/1499 RP - 16/3-2 Well (Norway)  
Sedimentological study of Jurassic deposits.

R. CUSSEY : 7/1517 - 16/3-2 Well (Norway)  
Sedimentological study of the Tertiary deposits.

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I L L U S T R A T I O N S

Plate 1 : Organic matter reflectance and fluorescence.

Plate 2 : Upper Jurassic. Characterization of the syngenetical hydrocarbons.

Plate 3 : Organic matter study. Synthesis of results.

This study was carried out to appreciate the source-rock quality and estimate the state of maturation of the organic matter between 1000 m and the well bottom, in particular at the level of the Hot Shales of the Upper Jurassic (Portlandian and Kimmeridgian , 1955 to 1975 m).

It takes into account observations of the kerogen, palynofacies and thermal alteration index, carried out on the palynological slides\*.

All the results are summarized on Plate 3, "Synthesis of results".

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\* by J.F RAYNAUD, in 2035 n° 6/27 N from april 1976, and by J. DUCAZEAUX, in GEO/LAB Bss n° 6/1418 R from december 1976.

1 - ORGANIC COMPOUND CONTENTS

The measurements were carried out on 9 shale sidewall cores below 1350 m. Apart from two samples, for which the extractable organic matter contents were measured (EOM, in ppm of rock), only the insoluble organic carbon contents were measured (IOC, in weight % of rock), due to the small quantities of rock available.

The results are given below (see also Pl. 3) :

		IOC	EOM
Paleocene	1360 m	1.50	
Cenomanian (?)	1744	1.00	
Albian	1811	0.85	
Aptian	1894	0.90	
Upper Jurassic	1959	7.90	
"	1961	5.30	3300
"	1967	4.20	2800
"	1969	5.40	
"	1973	8.00	

1955 } Portl.  
 1975 } Kimm.

They mainly show that the IOC of the Upper Jurassic shales (Hot Shales) are high : the average is approximately 6%.

The two EOM are relatively low : they show that about 5 % of the total organic matter of the Hot Shales is extractable.

2 - ORGANIC FACIES

2.1 - In transmitted light : Palynofacies (see Pl. 3)

Between 1230 and 1400 m (Middle to Low Eocene), the palynofacies are made up of a mixture, in variable proportions, of amorphous organic matter, opaque woody fragments and coals ; amorphous organic matter is predominant at 1310 and 1400 m.

Between 1410 and 1480 m (Danian), the palynofacies are mainly made up of coals associated with opaque woody fragments and black powdery matter.

Between 1490 and 1650 m (Upper Cretaceous), there is very little abundant organic matter.

At 1660 m the palynofacies is made up of coals.

Between 1660 and 1890 m no sample has been studied.

At 1894 and 1934,5 m (Lower Cretaceous), the palynofacies are made up of ligneous and coaly particles.

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At 1959 and 1973 m (Portlandian-Kimmeridgian), the palynofacies are mainly made up of amorphous organic matter.

At 2000 m (Kimmeridgian), the palynofacies is made up of amorphous matter, ligneous and coaly particles.

## 2.2 - In reflected white and UV light (see Pl. 1)

These observations were carried out on 16 samples (3 sidewall, 1 core and 12 cuttings), between 1000 m and 2010 m.

In reflected white light, within the lower reflectance population (.3 % to .8 %) humic coals and bitumens are present along the entire section ; however, a great part of them are probably unreliable, even in the sidewall core samples (reworking and mud contamination).

In reflected UV light, 2 main zones can be distinguished by their higher amount of fluorescent material :

- . at 1,000-1,100 m (the second being probably derived from the first), in the oligocene level, abundant algal matter and figured algae are present ; but due to the mixing of fresh (yellow) and evolved (dark orange) tasmanaceae, they are suspected to be partly or completely reworked matter.
- . the Upper Jurassic contains a certain amount of sapropelic matter : pure homogeneous alginite, marine algae (tasmanaceae) and sapropelic rock fragments with "bituminite" (from mainly algal origin) and algae.

The higher sample provided of this matter, is the sidewall core at 1961 m with a global amount in fluorescent material such as given by index 3 (on a total scale up to 5 for pure bogheads).

In addition, fluorescent reservoirs, due to traces of probably migrated hydrocarbons, have been observed at 2 levels :

- in small quantity at 1400 m in the Paleocene ;
- in a greater quantity with a few bitumens in the cutting sample, at 2010 m in argillaceous siltstones which are derived from Upper Cretaceous, according to the biostratigraphical survey (see § 5).

## 2.3 - Organic facies conclusion

The optical examinations in transmitted and reflected light point out that the organic content of the Upper Jurassic shales (Hot Shales) is mainly made up of amorphous partly sapropelic matter.

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### 3 - ORGANIC CATAGENESIS

#### 3.1 - In transmitted light : thermal alteration indices (see Pl. 3)

The thermal alteration indices estimated in the lower part of the well, between 1894 m and 2000 m, are 2 to 2.5.

#### 3.2 - In reflected white and UV light : reflectance and fluorescence (see Pl. 1 and 3)

The main results are not very good, due to the great dispersion of reflectance measurements and fluorescence observations, which attest to a large mixing of organic matter populations.

These populations consist of :

- reworked coal and bitumen particles with a reflectance dispersed from 1% to 4% (and even to 5.5%, a value which is not mentioned on Pl. 1).

Their presence, and abundance, is probably related to the metamorphic paleozoic series which supplied the regional sedimentation for all the Jurassic to Tertiary period.

- one particular anisotropic coal or bitumen, very similar to a natural coke, perhaps heated by a volcanic dyke and localized around the depths 1600 to 1700 m. Its reflectance varies between 1 and 4 %.
- the likely reliable population (bitumen and vitrinite measurements) is scattered and a good interpretation of its results can be obtained only in the bottom 50 meters of the well, thanks to associated data of reflectance and fluorescence : abundant algae tasmanaceae in the 3 levels 1961-2000-2010, fluorescence colour yellow to pale orange attest to a low level of diagenesis, such as 0.5-0.6 % reflectance, in agreement with that which has been measured on bitumen and vitrinite in corresponding samples.

#### 3.3 - Chromatographical data

The chromatograms (Pl. 2) of the syngenetical hydrocarbons from Hot Shales (5 samples between 1959 and 1973 m) show that these hydrocarbons are at a low stage of maturation. This is mainly indicated by the low n-alkane content of the thermovaporised and the saturated fractions ( $<13\%$  and  $<5\%$  respectively), the low X1 and X2\* ratios (lower than 1) and the higher isoprenoid contents in comparison to n-alkane contents (pristane/nC 17  $>1.5$  and phytane /nC 18  $>2$ ).

#### 3.4 - Organic catagenesis conclusions

All the data mentioned above shows that the maturation degree of the organic matter is low down to the well bottom, lower than the required for substantial oil generation

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\* X1 : nC6/Methylcyclopentane  
X2 : nC7/Dimethylcyclopentane

#### 4 - SOURCE ROCK CHARACTERISTICS

##### . Paleocene and Cretaceous, between 1350 and 1955 m

The thermovaporisations carried out on 3 sidewall cores (at 1360, 1744 and 1894 m) contain no hydrocarbon : this suggests a bad petrolenn quality, in spite of the low kerogen maturation .

The pyrolysis undertaken on the 1360 sample confirms that the initial hydrocarbon potential at this level is very low : the content in organic carbon degraded by this process is about 0.05 % of the rock, i.e. 5 % of the kerogen.

##### . Upper Jurassic. Portlandian and Kimmeridgian brown shales, between 1955 and 1975 m (see the chromatograms on Pl. 2)

The thermovaporisations undertaken on the 5 sidewall cores studied show variable contents in syngenetical hydrocarbons (between 200 and 800 ppm), relatively high in view of the low maturation.

The composition of two extracts (1961 and 1965 m) mainly shows a low percentage of hydrocarbons (< 30 %) ; the saturates/aromatics ratios are 1 and 0.6, but the highest value, at 1961 m, is probably due to contamination (see below).

The saturates chromatograms\* are characterized by a naphthenic development, some irregular n-alkane contents, abundant isoprenoids in comparison to n-alkanes and pristane/phytane ratios between 0.7 and 1 ; the chromatogram of the 1961 sample shows a high development of n-alkanes between C 22 and C 30 (maximum at C 25) suggesting contamination by heavy parafins.

The pyrolyses performed on 3 sidewall cores (at 1959, 1969 and 1973 m) show that the kerogen quality is good : the thermodegraded organic carbon contents are between 1.75 to 3 % of the rock, i.e. 35 % of the kerogen.

As the maturation is low, the addition of the contents in extractable and degradable organic carbon indicates that initially about 40 % of the organic matter may produce oil. Knowing that the average content in organic matter of the Upper Jurassic shales is about 6 % and their thickness about 20 m, their initial genetic potential is approximately  $1.5 \times 10^6$  T/km<sup>2</sup>.

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\* Chromatography (GPC) of the saturated fraction of the extract by chloroform (at 1961 and 1967 m) or of the extract by Hexane ("minianalyses" at 1959, 1969 and 1973).



5 - ATTEMPT TO CHARACTERIZE MIGRATED HYDROCARBONS

*low conc -*  
↑  
The examinations in UV light of the Kimmeridgian coarse sandstones (core samples) showed neither fluorescence on rock nor fluorescence on extract, so these sandstones contain no significant quantity of hydrocarbons.

A minianalysis carried out on handpicked argillaceous siltstones from cuttings at 2010 m, in which some fluorescent hydrocarbon traces and bitumens were observed in reflected UV light (Pl. 1), showed only a few ppm of saturated hydrocarbons. These siltstones are caved from Upper Cretaceous according to a biostratigraphical survey.

6 - CONCLUSIONS

A good virtual source-rock interval is present in the 16/3-2 well between 1955 and 1975, i.e. the Hot Shales of the Upper Jurassic.

Its initial hydrocarbon potential is about  $1.5 \times 10^6$  T/km<sup>2</sup>; its maturation is lower than that corresponding to the main phase of oil generation.

REFLECTANCE  
(measurements)

FLUORESCENCE  
(Global amount estimated)

□ Vitrinite  
 ▨ Bitumens

■ Fluorescent macerals  
 ▩ Hydrocarbon traces in reservoirs

MP Mud products

R E F L E C T A N C E

F L U O R E S C E N C E

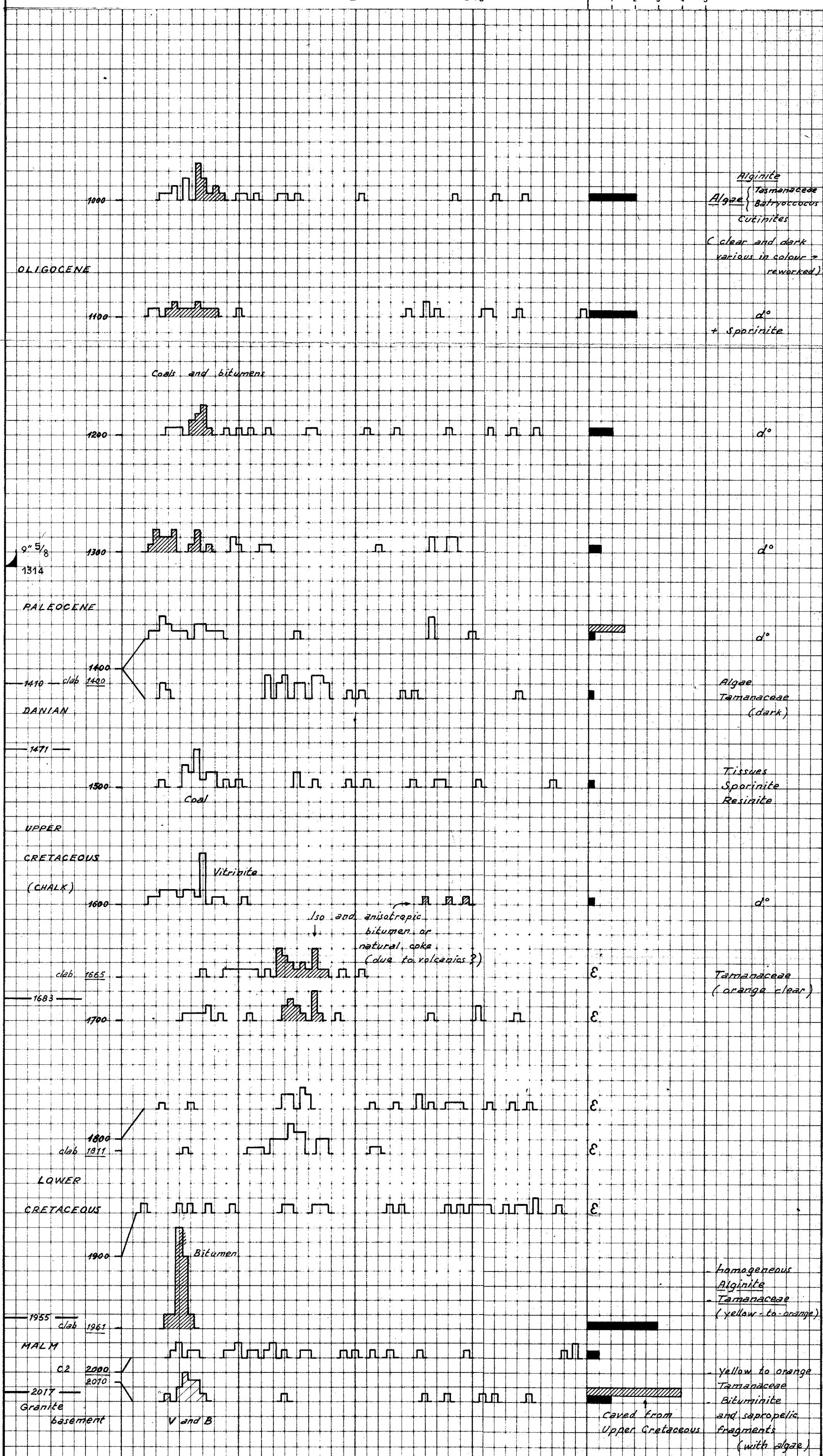
VITRINITE  
BITUMEN

GLOBAL  
ESTIMATED

ORGANIC MATTER  
NATURE

0                      1                      2                      3%

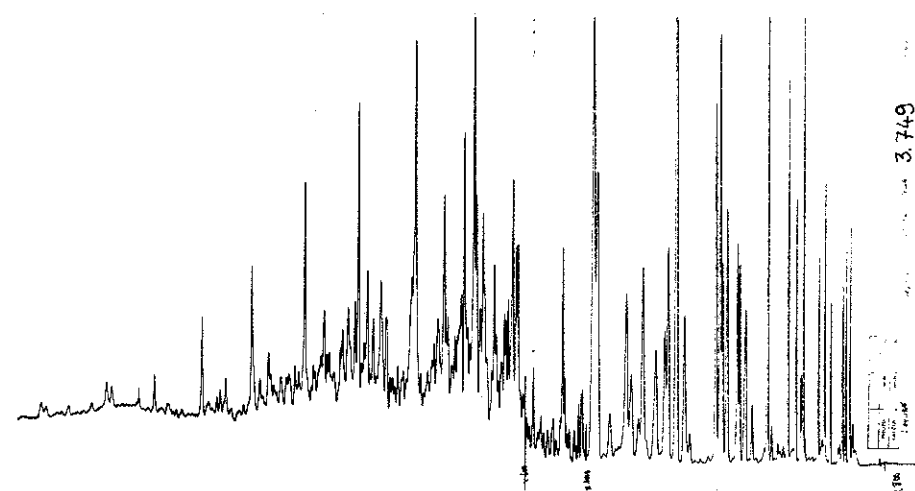
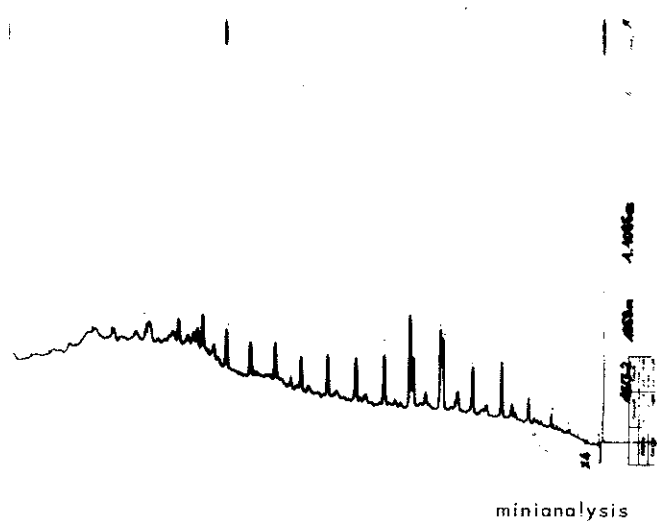
1    2    3    4    5



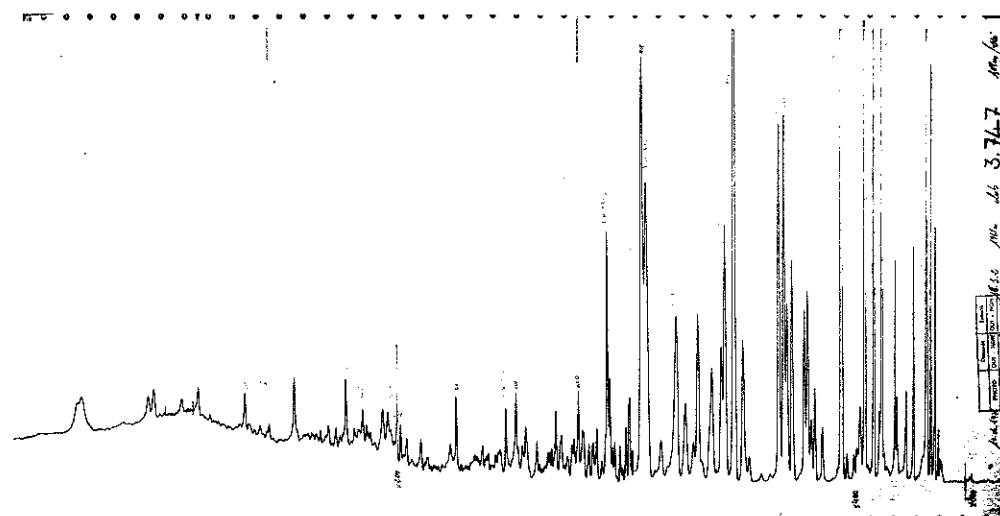
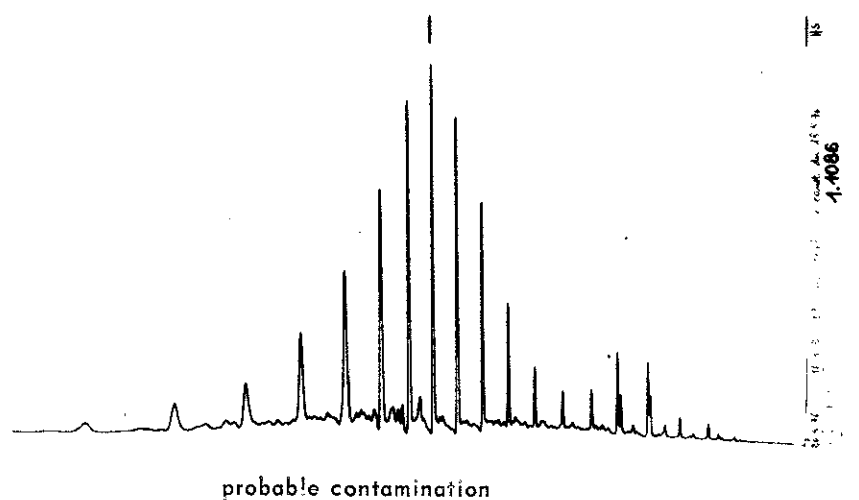
SATURATED HC

THERMOVAPORIZED HC

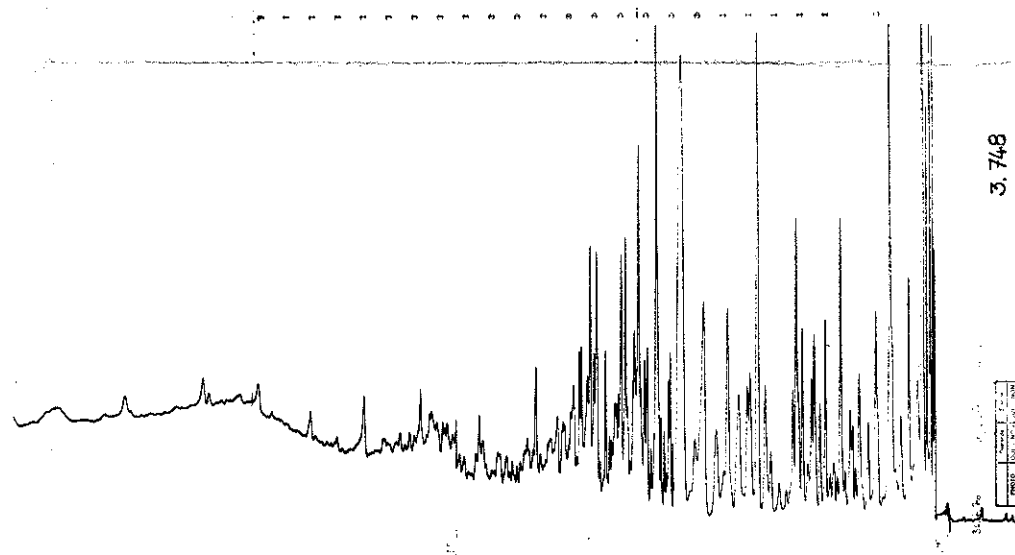
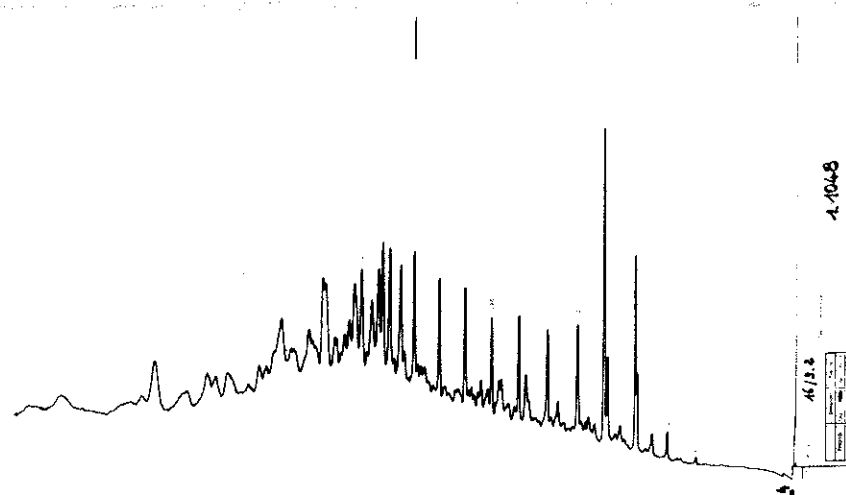
1959 m  
COI = 79 %



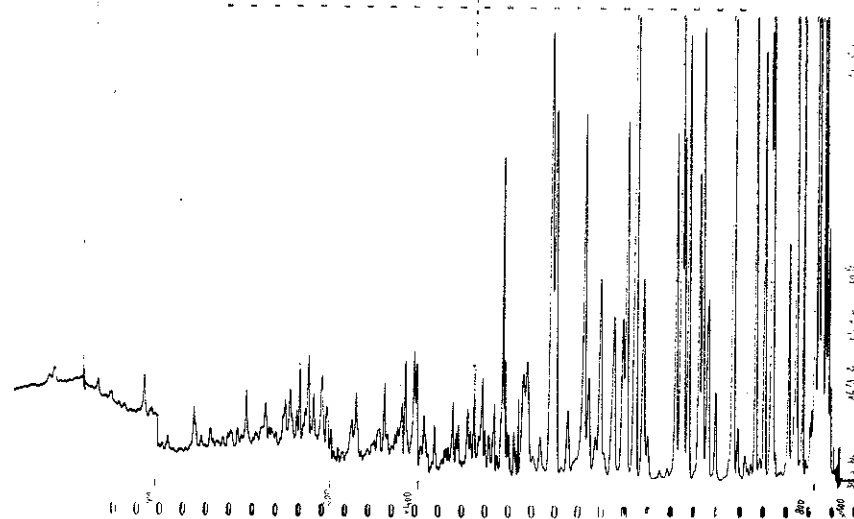
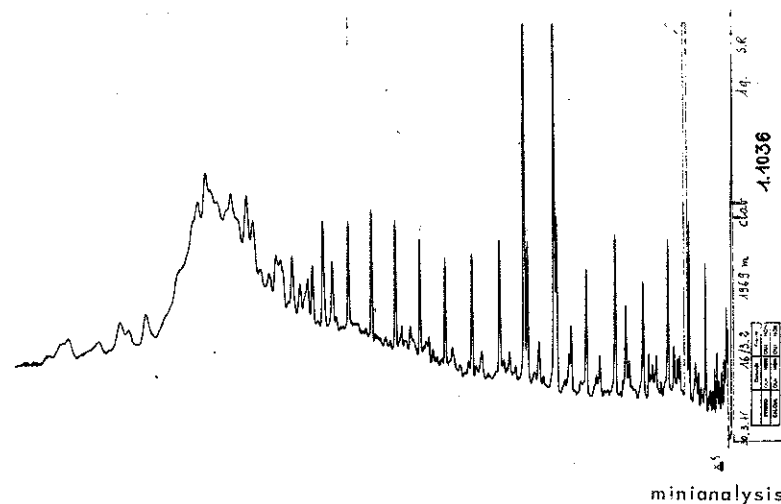
1961 m  
COI = 53 %  
MOE 3000 ppm



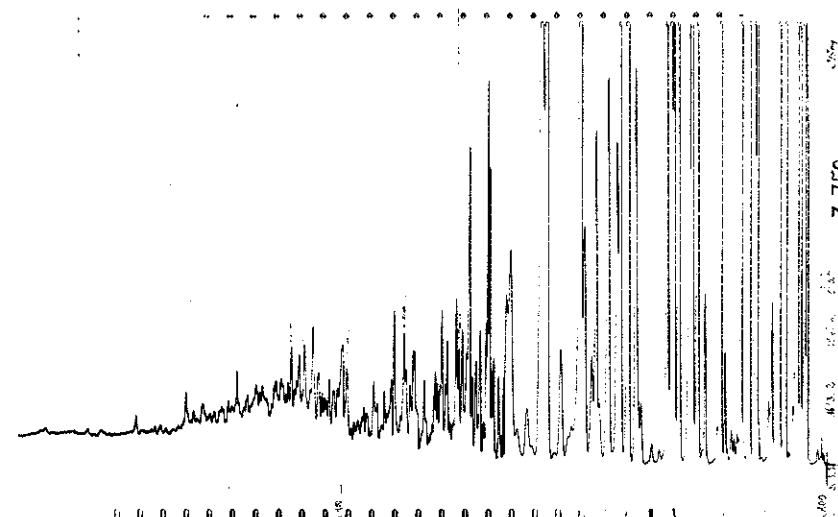
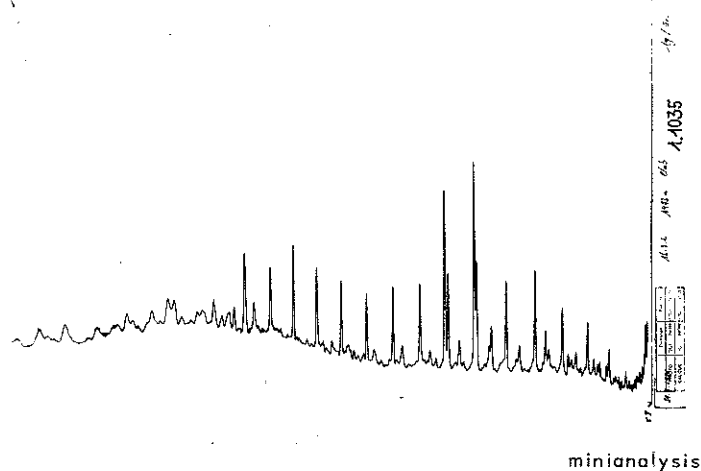
1967 m  
COI = 42 %  
MOE 2700 ppm



1969 m  
COI = 54 %



1973 m  
COI = 8 %



elf	Secteur	OFF SHORE NORWAY	PETRONORD
	Opérateur	NORSK HYDRO-ELF	
	Permis de Concession	007	
<b>16/3.2 WELL</b>			
<b>UPPER JURASSIC</b>			
CHARACTERIZATION OF THE SYNGENETICAL HYDROCARBONS (Sidewall core samples)			
OFF ENTREPRISE DE RECHERCHES ET D'ACTIVITES PETROLIERES			Date: Sep. 1977
DIRECTION EXPLORATION			PL. 2
LABORATOIRE			N°classif C 1855

## Synthesis of results

